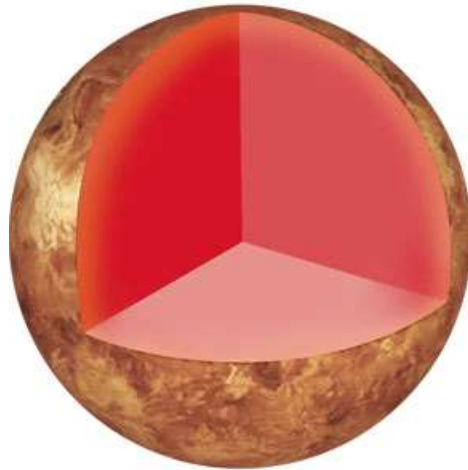


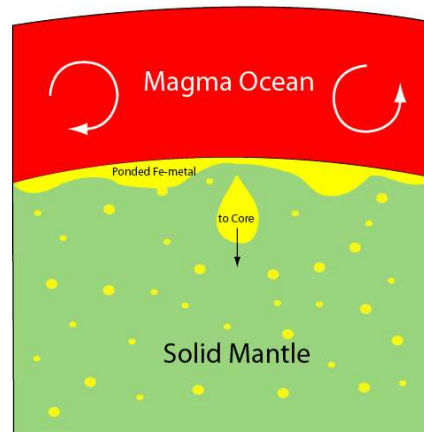
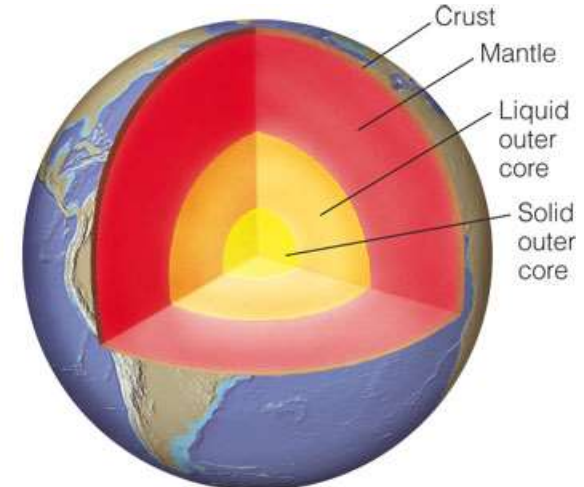
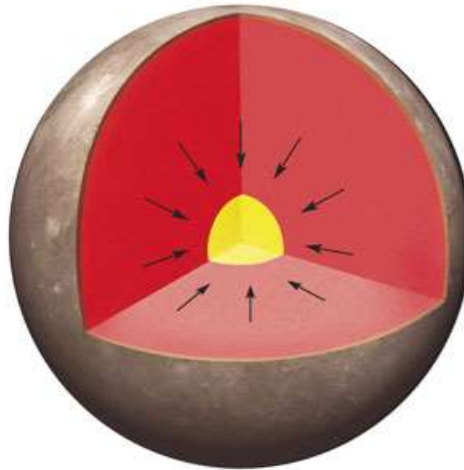
# Lecture 2: Before we get to PLATE TECTONICS.....

## *A brief accounting of Earth's composition, interior structure, and time periods*

Proto-Earth was almost entirely molten **Time** → Iron-nickel core surrounded by silica-rich mantle



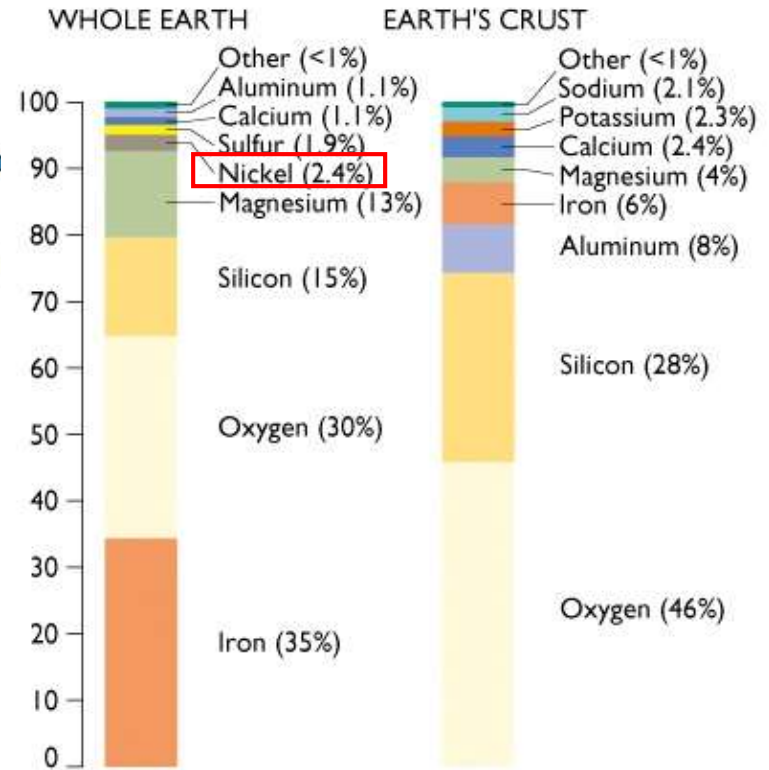
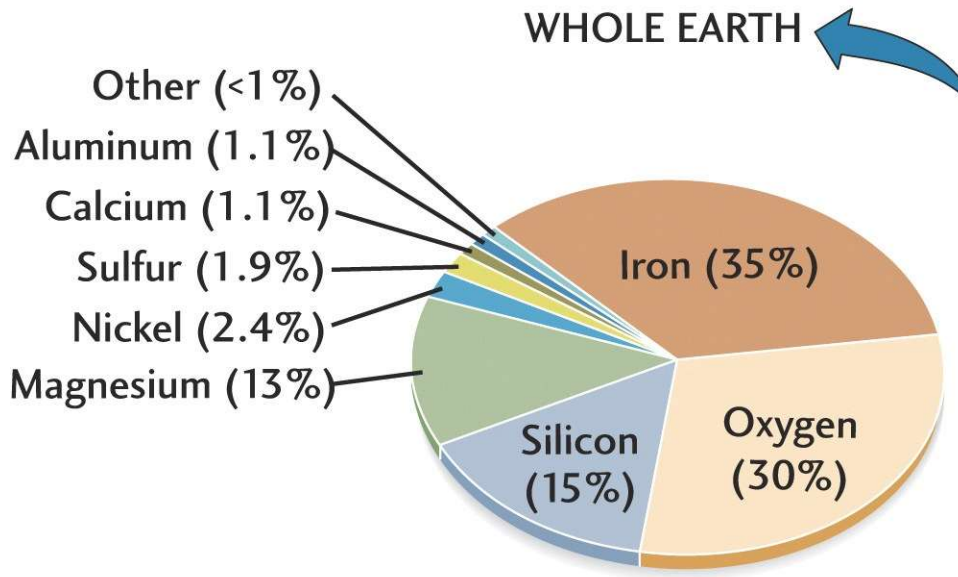
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Separation, or *differentiation*, of iron from lighter elements results in stratification, that can happen within hundreds of years in a magma ocean

[www.geodynamics.rice.edu/tobias.hoeink/index.php?page=research](http://www.geodynamics.rice.edu/tobias.hoeink/index.php?page=research)

# Earth's composition



# The 26 naturally occurring elements formed by stellar nucleosynthesis

Common in: all crust  continental crust  oceanic crust and mantle 

hydrogen 1 <b>H</b> 1.0079												helium 2 <b>He</b> 4.0026													
lithium 3 <b>Li</b> 6.941		beryllium 4 <b>Be</b> 9.0122												boron 5 <b>B</b> 10.811		carbon 6 <b>C</b> 12.011	nitrogen 7 <b>N</b> 14.007	oxygen 8 <b>O</b> 15.999	fluorine 9 <b>F</b> 18.998	neon 10 <b>Ne</b> 20.180					
sodium 11 <b>Na</b> 22.990		magnesium 12 <b>Mg</b> 24.305												aluminum 13 <b>Al</b> 26.982	silicon 14 <b>Si</b> 28.086	phosphorus 15 <b>P</b> 30.974	sulfur 16 <b>S</b> 32.065	chlorine 17 <b>Cl</b> 35.453	argon 18 <b>Ar</b> 39.948						
potassium 19 <b>K</b> 39.098		calcium 20 <b>Ca</b> 40.078		scandium 21 <b>Sc</b> 44.956	titanium 22 <b>Ti</b> 47.867	vanadium 23 <b>V</b> 50.942	chromium 24 <b>Cr</b> 51.996	manganese 25 <b>Mn</b> 54.938	iron 26 <b>Fe</b> 55.845	cobalt 27 <b>Co</b> 58.933	nickel 28 <b>Ni</b> 58.693	copper 29 <b>Cu</b> 63.546	zinc 30 <b>Zn</b> 65.39	gallium 31 <b>Ga</b> 69.723	germanium 32 <b>Ge</b> 72.61	arsenic 33 <b>As</b> 74.922	selenium 34 <b>Se</b> 78.96	bromine 35 <b>Br</b> 79.904	krypton 36 <b>Kr</b> 83.80						
rubidium 37 <b>Rb</b> 85.468		strontium 38 <b>Sr</b> 87.62		yttrium 39 <b>Y</b> 88.906	zirconium 40 <b>Zr</b> 91.224	niobium 41 <b>Nb</b> 92.906	molybdenum 42 <b>Mo</b> 95.94	technetium 43 <b>Tc</b> [98]	ruthenium 44 <b>Ru</b> 101.07	rhodium 45 <b>Rh</b> 102.91	palladium 46 <b>Pd</b> 106.42	silver 47 <b>Ag</b> 107.87	cadmium 48 <b>Cd</b> 112.41	indium 49 <b>In</b> 114.82	tin 50 <b>Sn</b> 118.71	antimony 51 <b>Sb</b> 121.76	tellurium 52 <b>Te</b> 127.60	iodine 53 <b>I</b> 126.90	xenon 54 <b>Xe</b> 131.29						
caesium 55 <b>Cs</b> 132.91		barium 56 <b>Ba</b> 137.33		lanthanum 57 <b>La</b> 138.91	hafnium 72 <b>Hf</b> 178.49	tantalum 73 <b>Ta</b> 180.95	tungsten 74 <b>W</b> 183.84	rhenium 75 <b>Re</b> 186.21	osmium 76 <b>Os</b> 190.23	iridium 77 <b>Ir</b> 192.22	platinum 78 <b>Pt</b> 195.08	gold 79 <b>Au</b> 196.97	mercury 80 <b>Hg</b> 200.59	thallium 81 <b>Tl</b> 204.38	lead 82 <b>Pb</b> 207.2	bismuth 83 <b>Bi</b> 208.98	polonium 84 <b>Po</b> [209]	astatine 85 <b>At</b> [210]	radon 86 <b>Rn</b> [222]						
francium 87 <b>Fr</b> [223]		radium 88 <b>Ra</b> [226]		actinoids 89-102 <b>**</b>	lawrencium 103 <b>Lr</b> [262]	rutherfordium 104 <b>Rf</b> [261]	dubnium 105 <b>Db</b> [262]	seaborgium 106 <b>Sg</b> [266]	bohrium 107 <b>Bh</b> [264]	hassium 108 <b>Hs</b> [269]	meitnerium 109 <b>Mt</b> [268]	unnilium 110 <b>Uun</b> [271]	unununium 111 <b>Uuu</b> [272]	ununbium 112 <b>Uub</b> [277]	ununquadium 114 <b>Uuq</b> [289]										

Key:

element name
atomic number
symbol
atomic weight (mean relative mass)

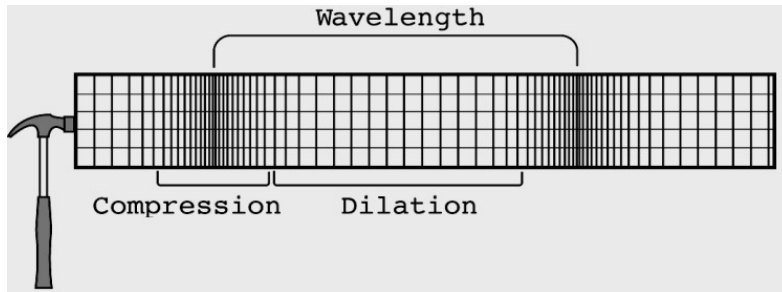
\*lanthanoids

\*\*actinoids

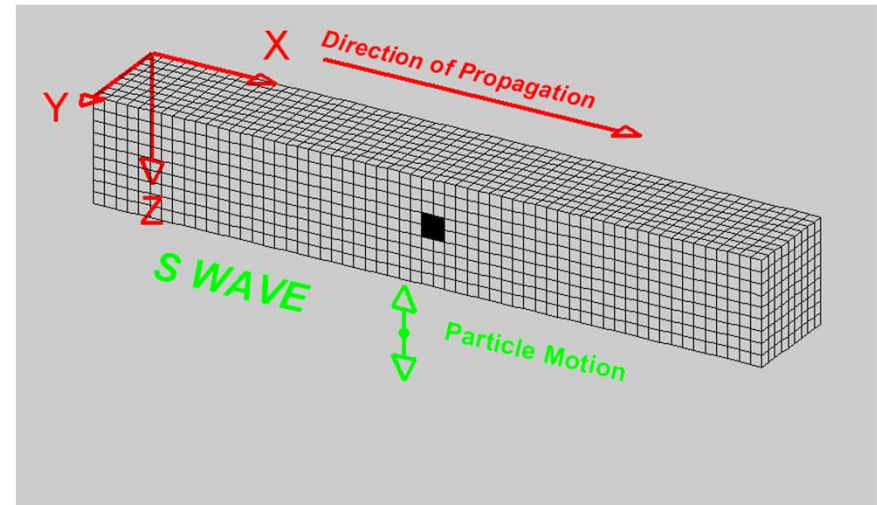
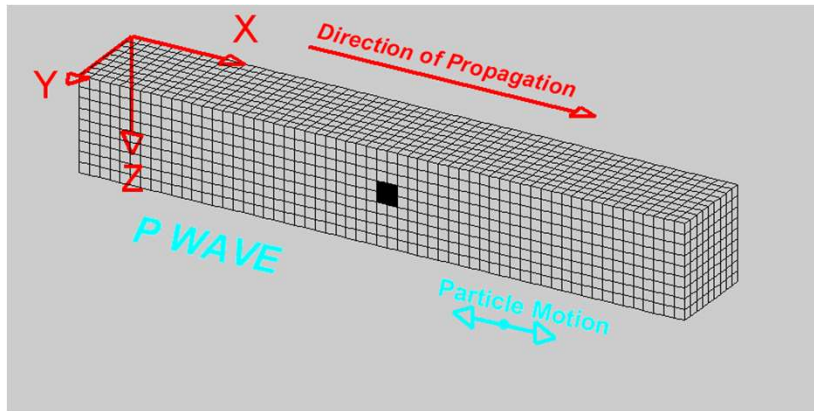
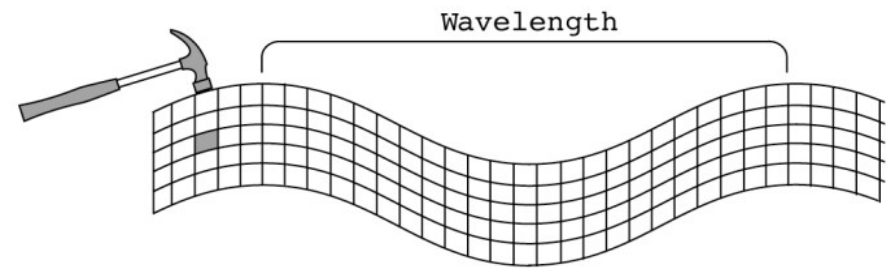
lanthanum 57 <b>La</b> 138.91	cerium 58 <b>Ce</b> 140.12	praseodymium 59 <b>Pr</b> 140.91	neodymium 60 <b>Nd</b> 144.24	promethium 61 <b>Pm</b> [145]	samarium 62 <b>Sm</b> 150.36	europium 63 <b>Eu</b> 151.96	gadolinium 64 <b>Gd</b> 157.25	terbium 65 <b>Tb</b> 158.93	dysprosium 66 <b>Dy</b> 162.50	holmium 67 <b>Ho</b> 164.93	erbium 68 <b>Er</b> 167.26	thulium 69 <b>Tm</b> 168.93	ytterbium 70 <b>Yb</b> 173.04
actinium 89 <b>Ac</b> [227]	thorium 90 <b>Th</b> 232.04	protactinium 91 <b>Pa</b> 231.04	uranium 92 <b>U</b> 238.03	neptunium 93 <b>Np</b> [237]	plutonium 94 <b>Pu</b> [244]	americium 95 <b>Am</b> [243]	curium 96 <b>Cm</b> [247]	berkelium 97 <b>Bk</b> [247]	californium 98 <b>Cf</b> [251]	einsteinium 99 <b>Es</b> [252]	fermium 100 <b>Fm</b> [257]	mendelevium 101 <b>Md</b> [258]	nobelium 102 <b>No</b> [259]

# EARTHQUAKES BODY WAVES

## P-waves

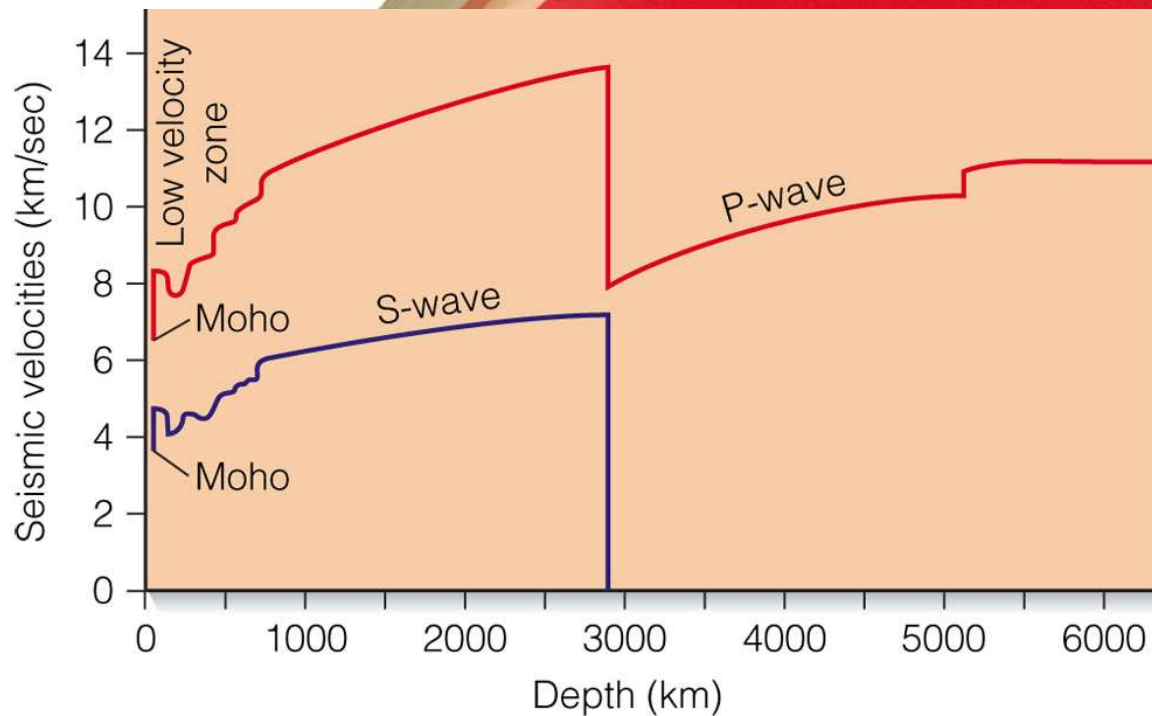
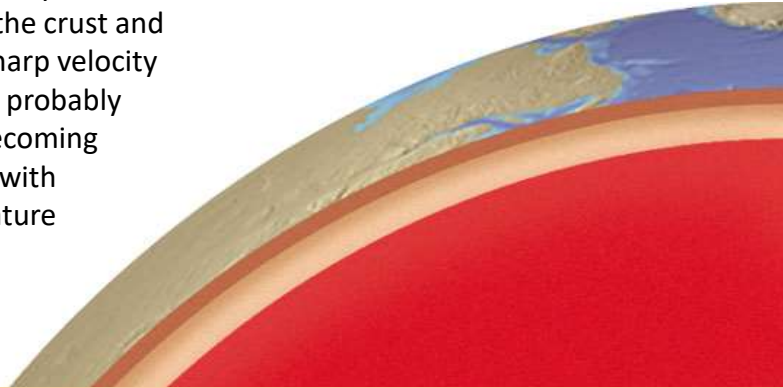


## S-waves



# Earth's crust, lithosphere, asthenosphere, mantle, and Moho

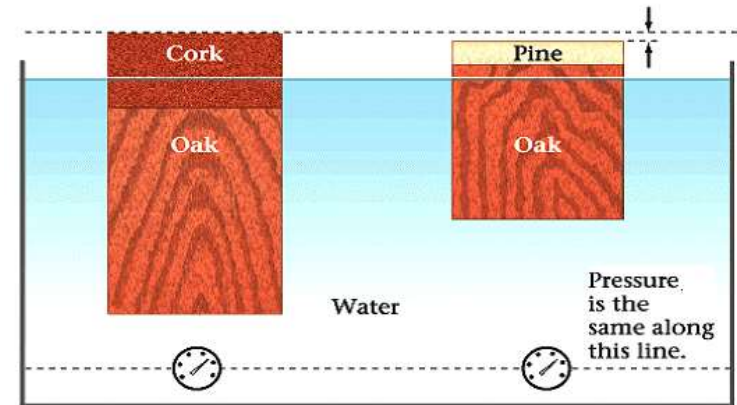
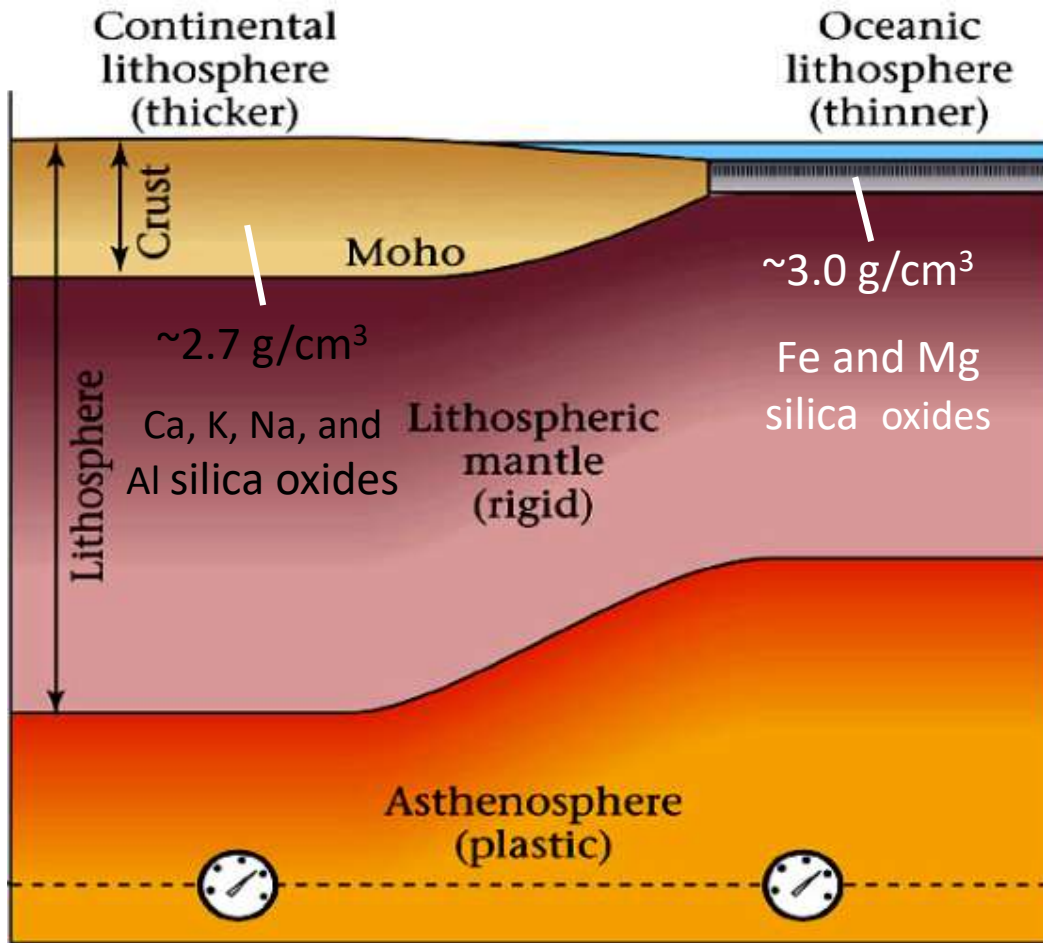
The Mohorovičić discontinuity, or **Moho**, is the boundary between the crust and the mantle defined by a sharp velocity increase in seismic waves, probably resulting from minerals becoming more dense and compact with increasing depth, temperature and pressure.



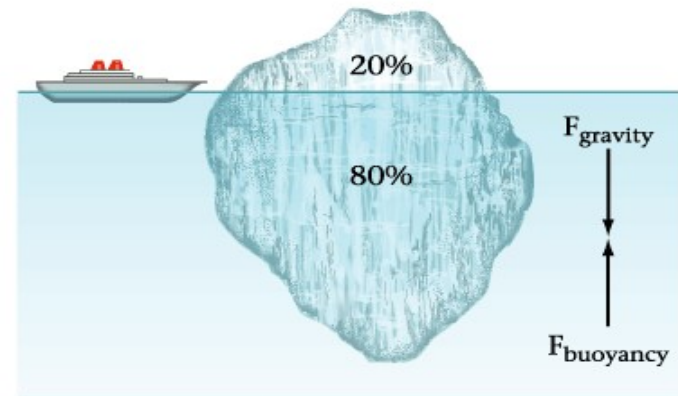
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- The lithosphere includes the crust and the uppermost mantle, which constitute the hard outer layer of the Earth that is rigid for very long periods of geologic time in which it deforms elastically and through brittle failure.
- The lithosphere is underlain by the asthenosphere, the weaker, part of the upper mantle that is hotter than the lithosphere, behaves more like plastic, deforms viscously and accommodates strain through plastic deformation.
- The lithosphere is broken into tectonic plates.

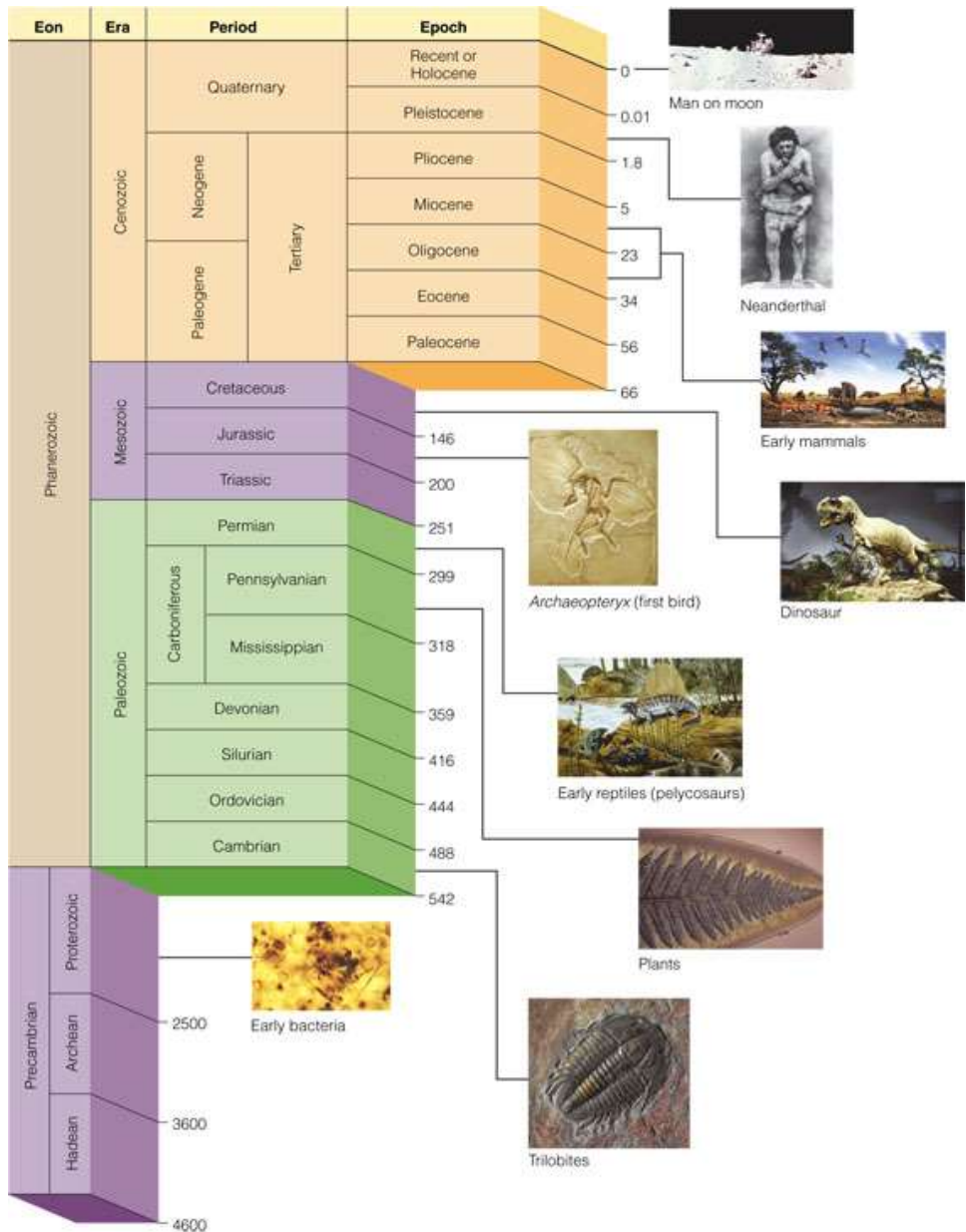
Oceanic crust is composed mostly of iron (Fe) and magnesium (Mg) silica (Si) oxides ( $O_n$ ) that are more dense and heavy than continental crust that has abundant sodium (Na), calcium (Ca), potassium (K), and Aluminum (Al) silica oxides



Continental crust has higher elevations and deeper roots than ocean crust because of gravity and buoyancy.



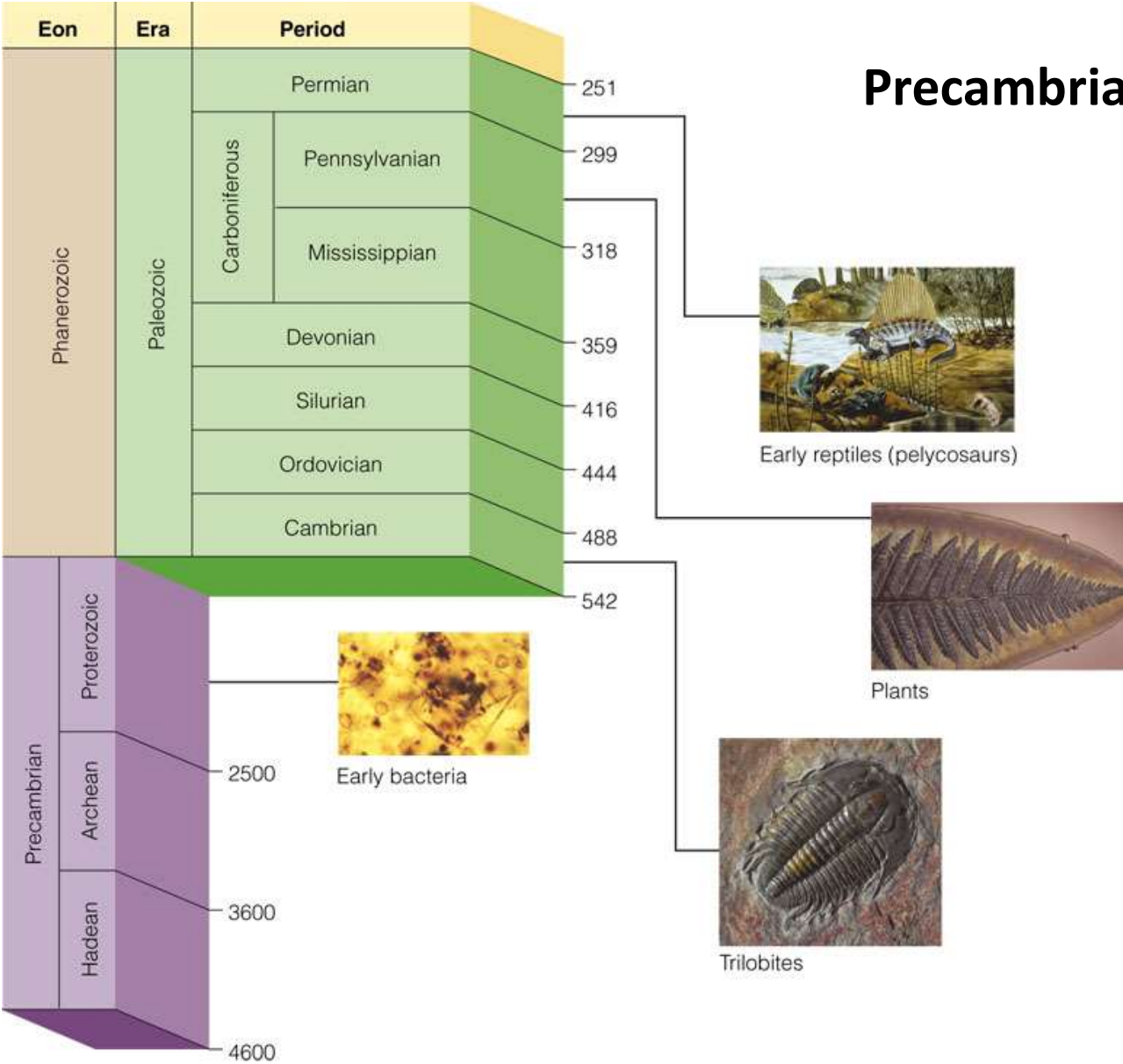
W. W. Norton. Modified from Sloss, NOAA



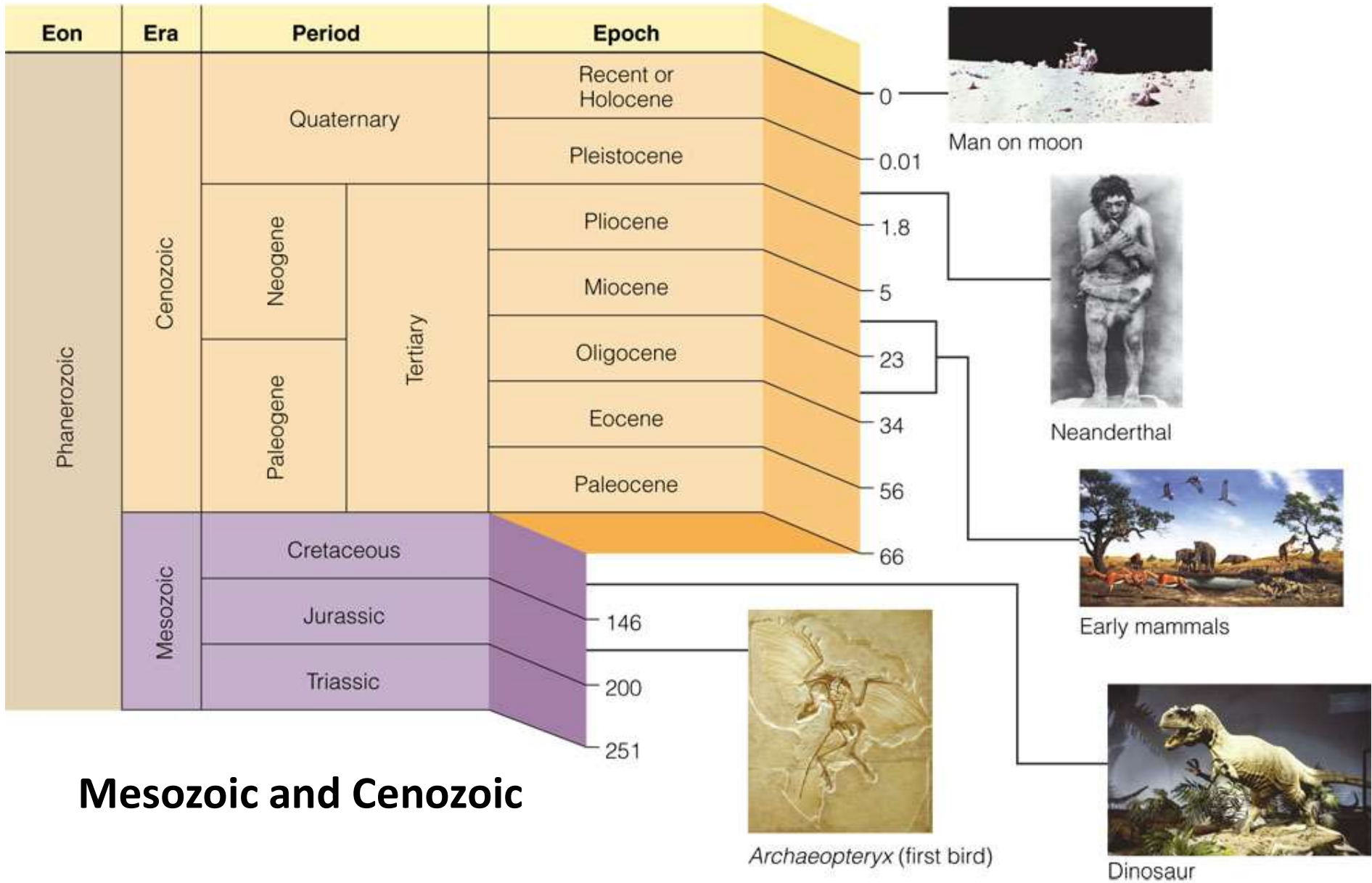
# Geologic time on Earth

- 4.6 Ba to the present
- Eon – billions to hundreds of millions
- Era - hundreds to tens of millions
- Period – tens of millions
- Epoch – tens of millions to tens of thousands

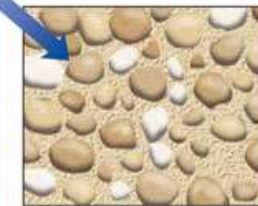
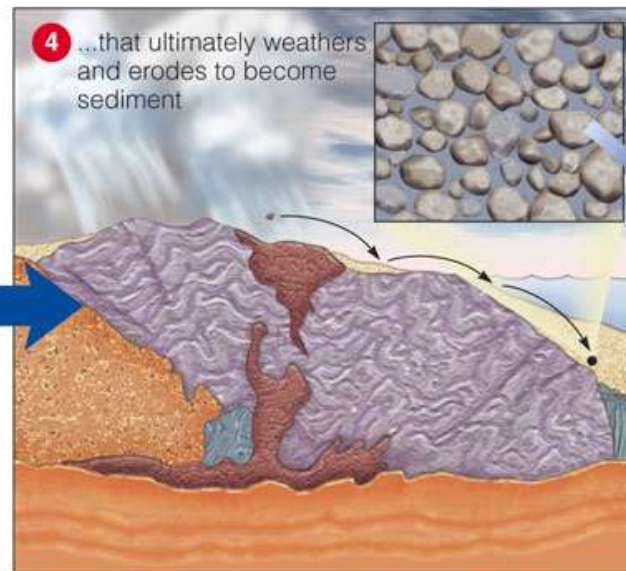
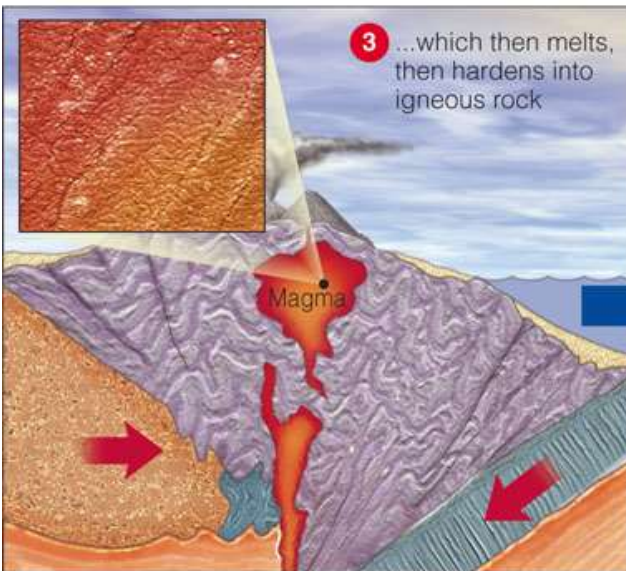
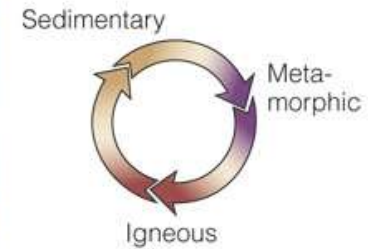
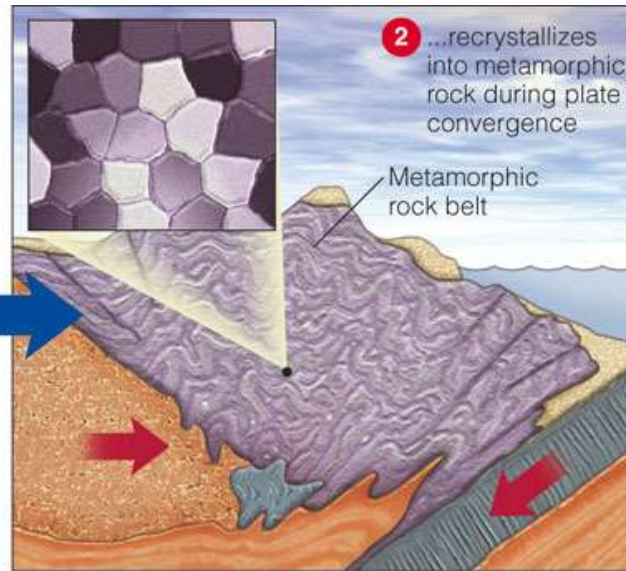
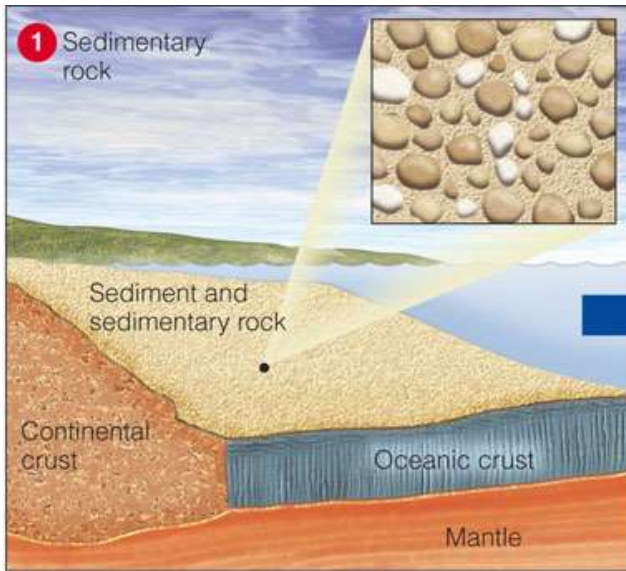
# Precambrian and Paleozoic







# Mesozoic and Cenozoic

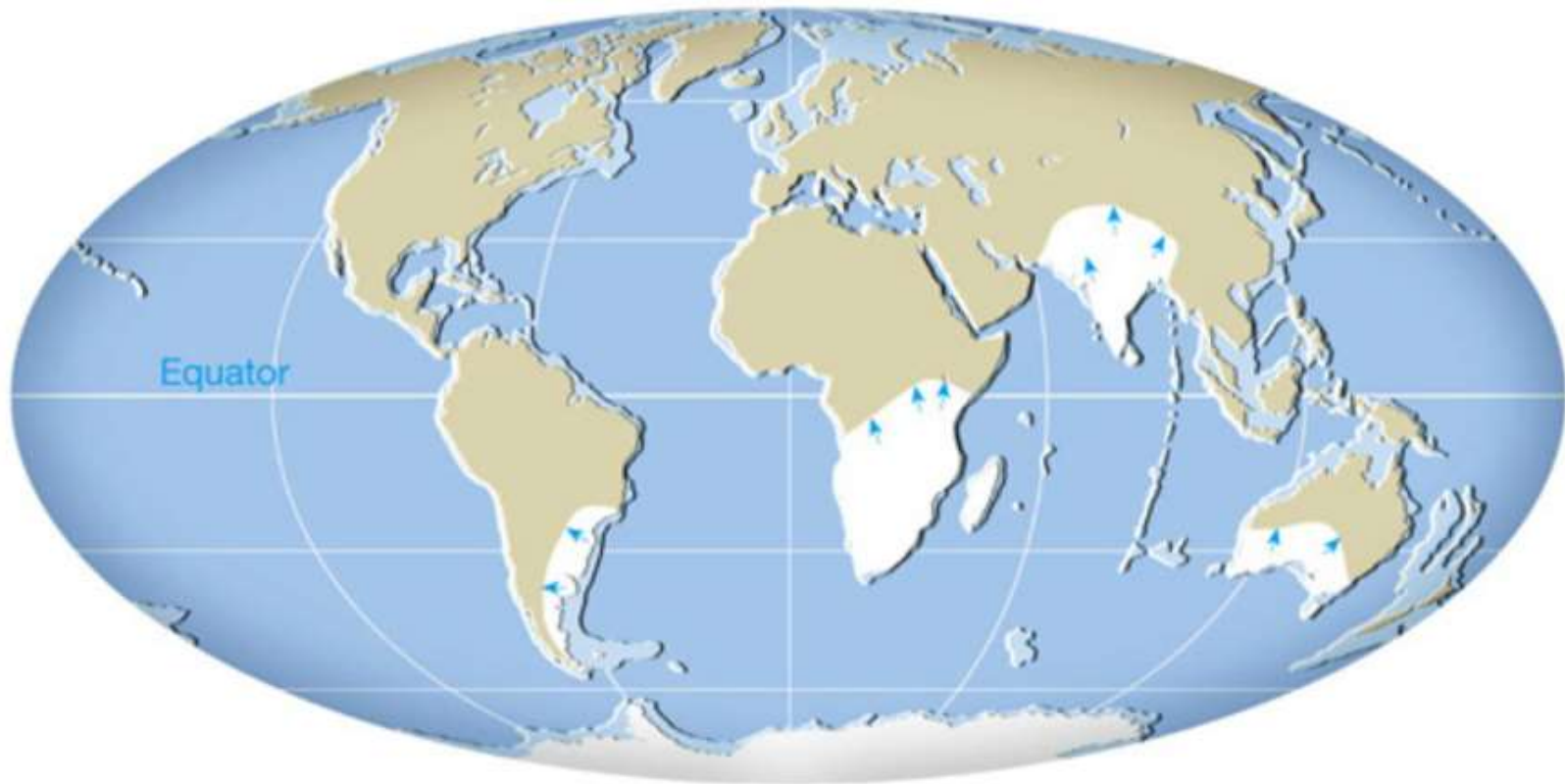


**5** ...which cements into hard rock. A continent has grown and the cycle starts all over again!

# The Rock Cycle

## CHAPTER 2 Plate Tectonics: A unifying theory

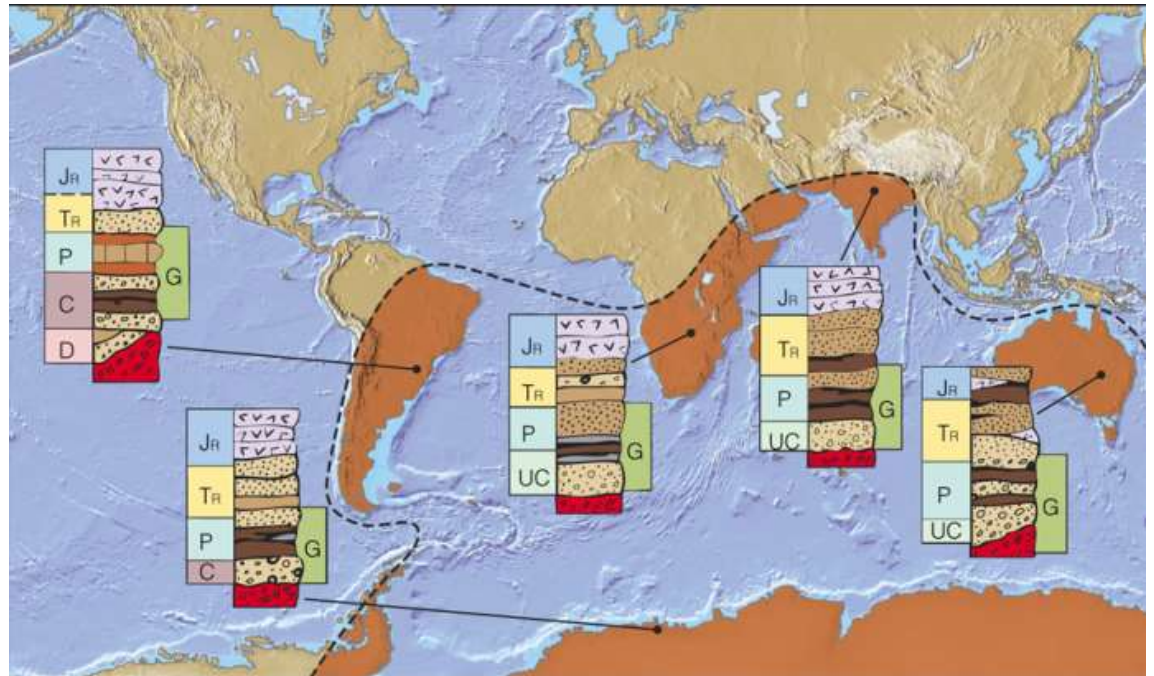
- The concept of continental movement was first suggested when it was noticed that Africa and South America had coastlines which appeared to be counterparts of one another, suggesting they may once have been joined and drifted apart.



- The hypothesis of continental drift is generally credited to Alfred Wegener, a German meteorologist who wrote the book, *The Origin of Continents*, first published in 1915.



- Wegener presented paleontological and geological evidence that the continents were once united into one supercontinent which he named Pangaea (Greek - “all land”).



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- His treatment and hypothesis were the most extensive and well developed at the time, but he could not provide a convincing mechanism to demonstrate how the continents could have moved, and his ideas were largely ignored.

- Wegener made mistakes identifying fossils, thought that the continents moved rapidly, up to meters per year, and thought that continents moved through oceanic crust, like ice-breaking ships.
- He disappeared in Greenland while trying to measure the rate of continental drift. The textbook reports that his presumed death was not widely mourned...



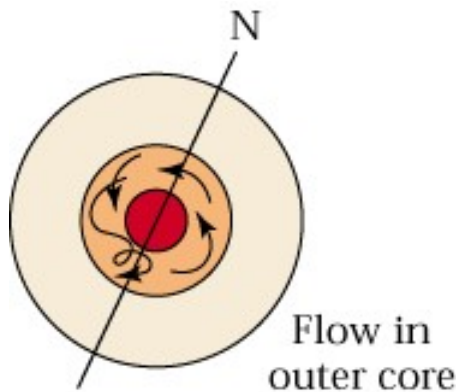
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***“Progress in science isn’t easy,  
and the arguments that  
ultimately drive science  
forward are often unpleasant.*”**

***In the case of Alfred Wegener,  
it is now clear that his insights  
outweighed his errors.”***

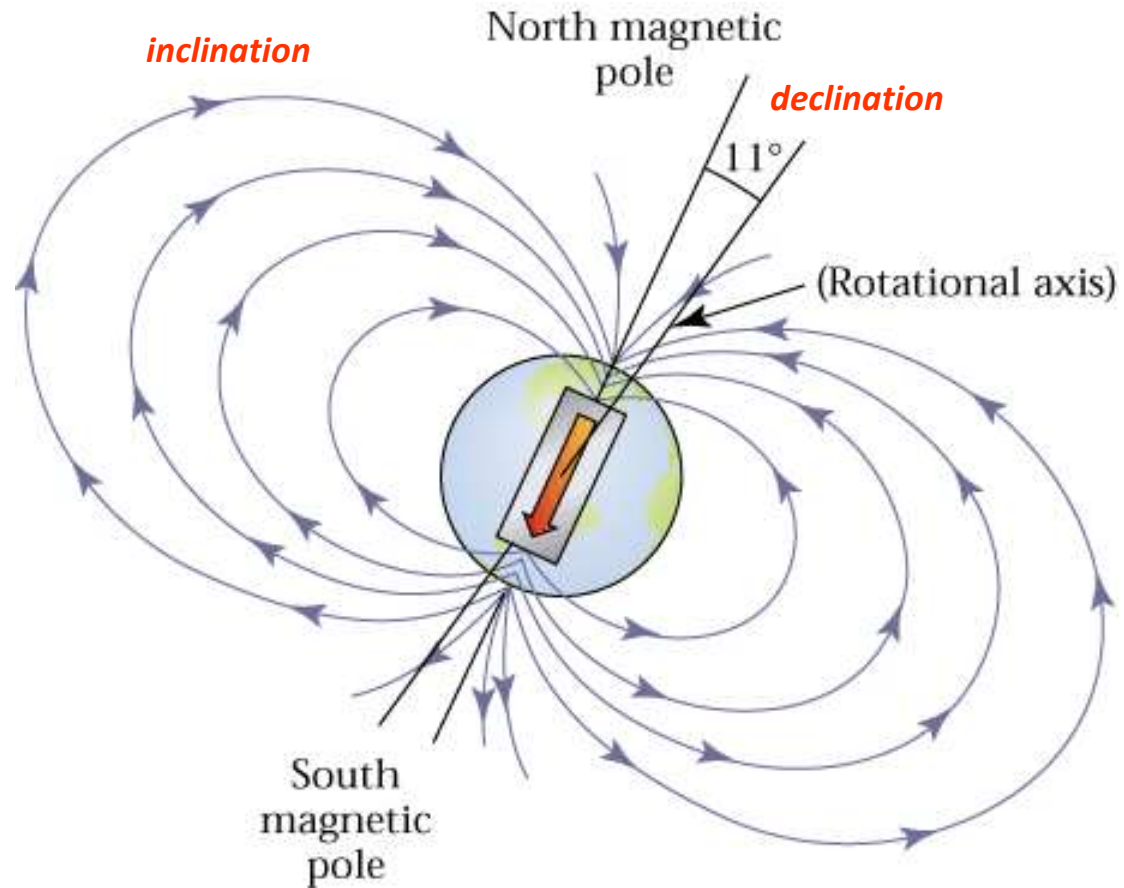
The 1950's brought a renewed interest to continental drift because of paleomagnetic studies.

**Paleomagnetism** is the remnant magnetism in ancient rocks recording the direction and intensity of the Earth's magnetic field.



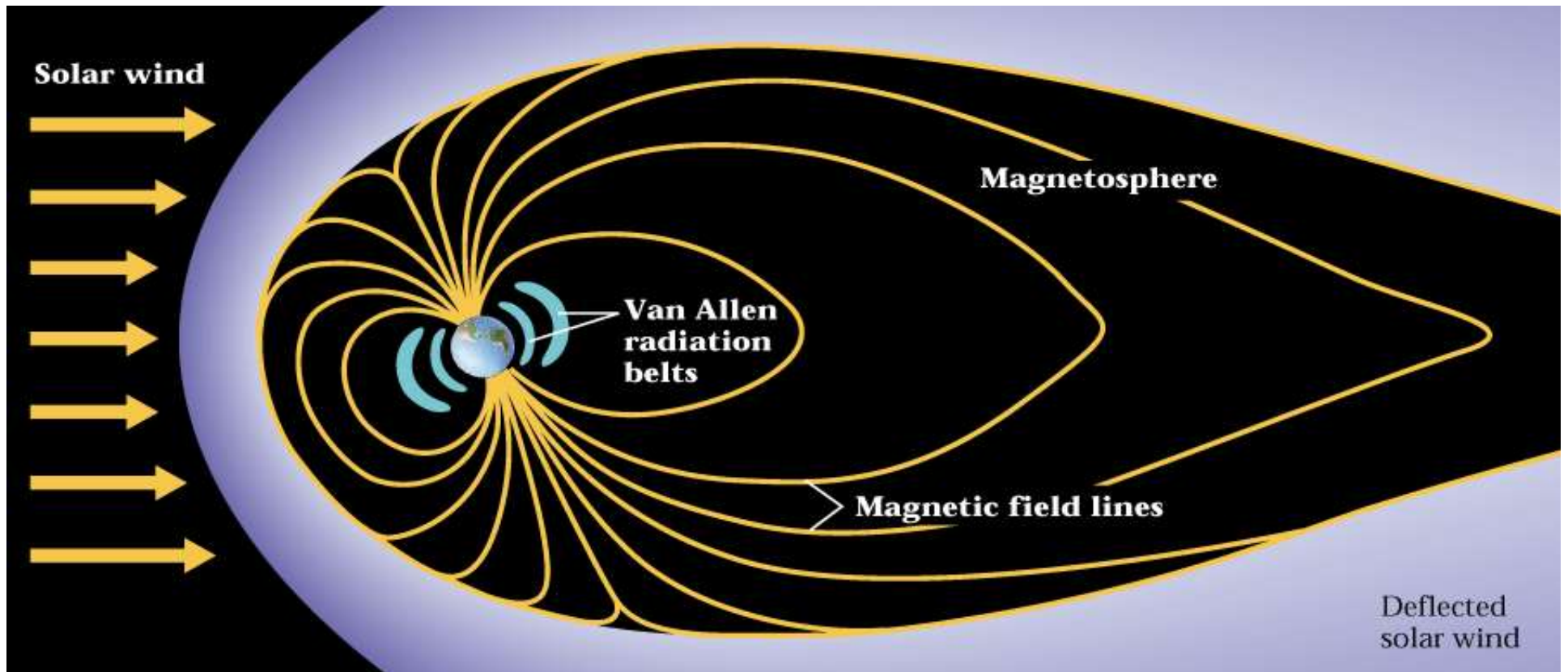
The Earth's magnetic field is generated by iron-rich, magnetic fluids in the outer core circulating around a spinning, solid-iron inner core.

This creates a "self-exciting dynamo".

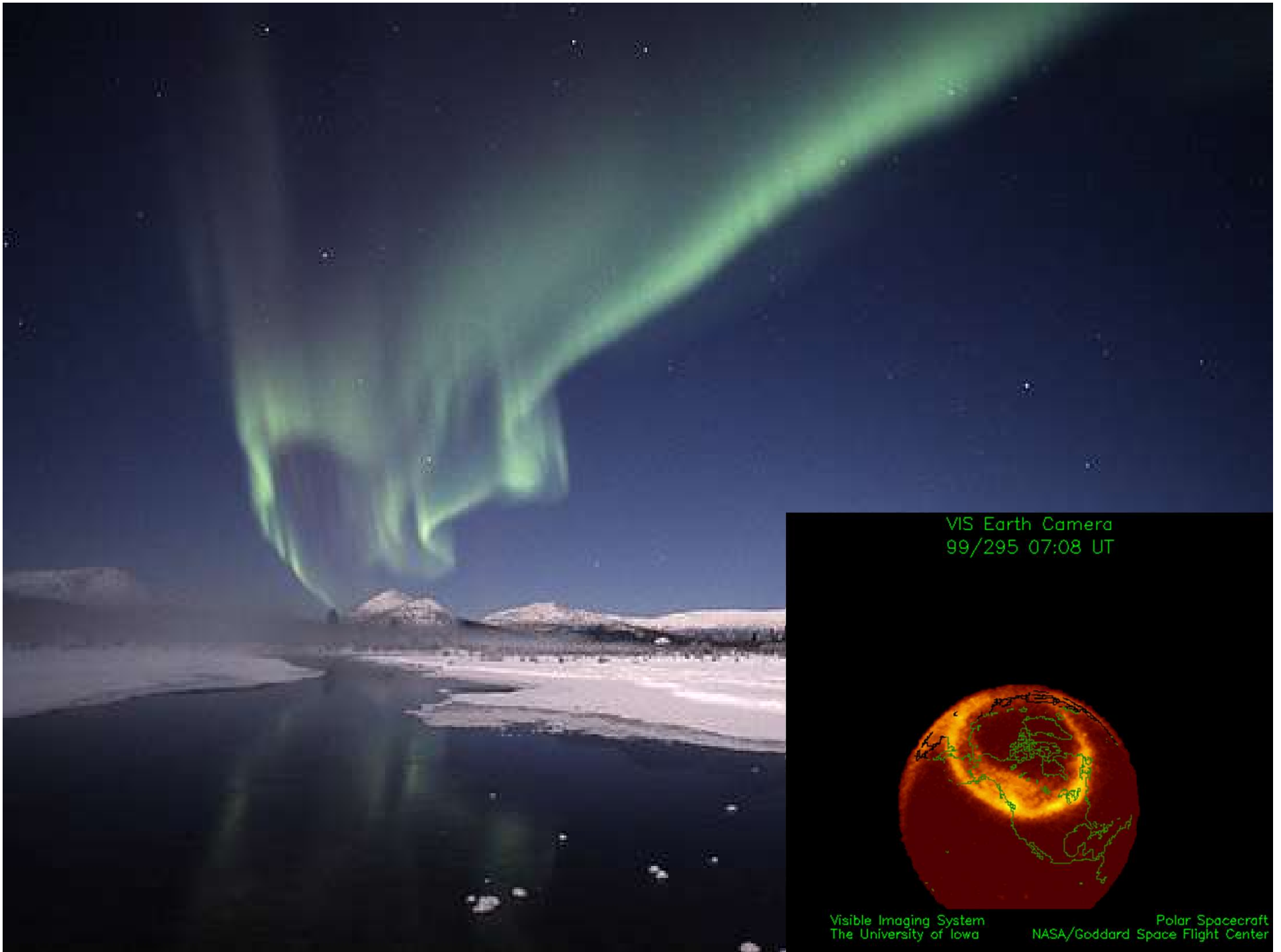


**Intensity of the magnetic field is greatest at poles and least at equator**

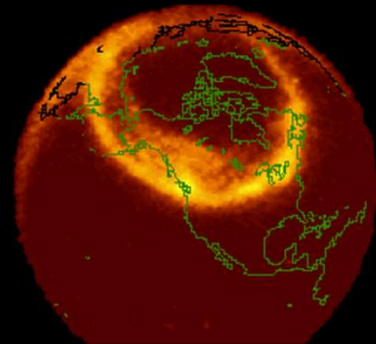
Our magnetic field protects us from UV radiation by deflecting low-energy charged particles *around* the Earth, but some particles become trapped in outer [Van Allen radiation belts](#).



Other high-energy particles are steered into vertical fields above N and S magnetic poles where they collide with ions in Earth's atmosphere and produce...



VIS Earth Camera  
99/295 07:08 UT

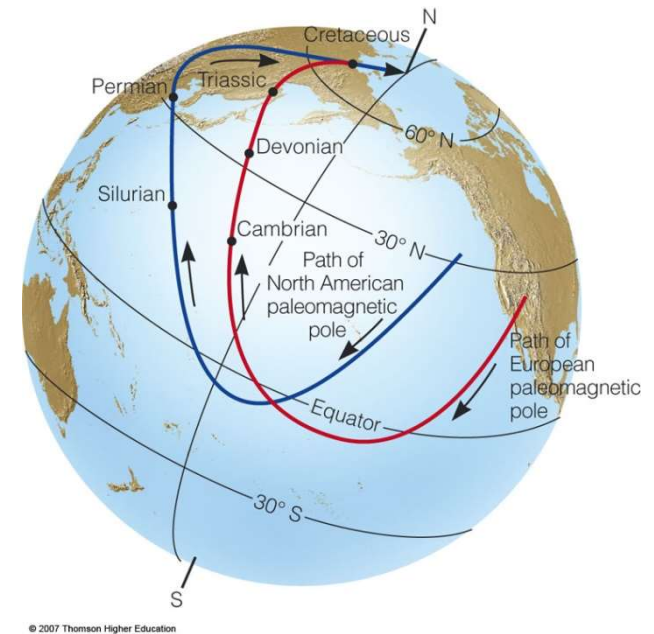
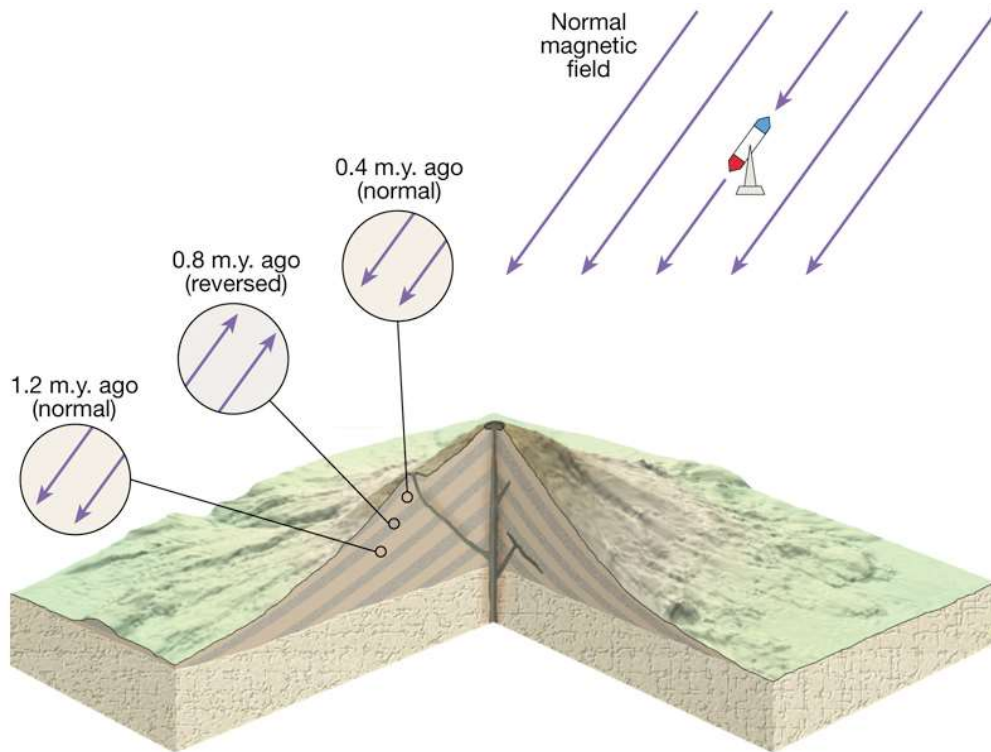


Visible Imaging System  
The University of Iowa

Polar Spacecraft  
NASA/Goddard Space Flight Center



- Studies of paleomagnetic poles from continental lava flows of different age show reversals of magnetic polarity.

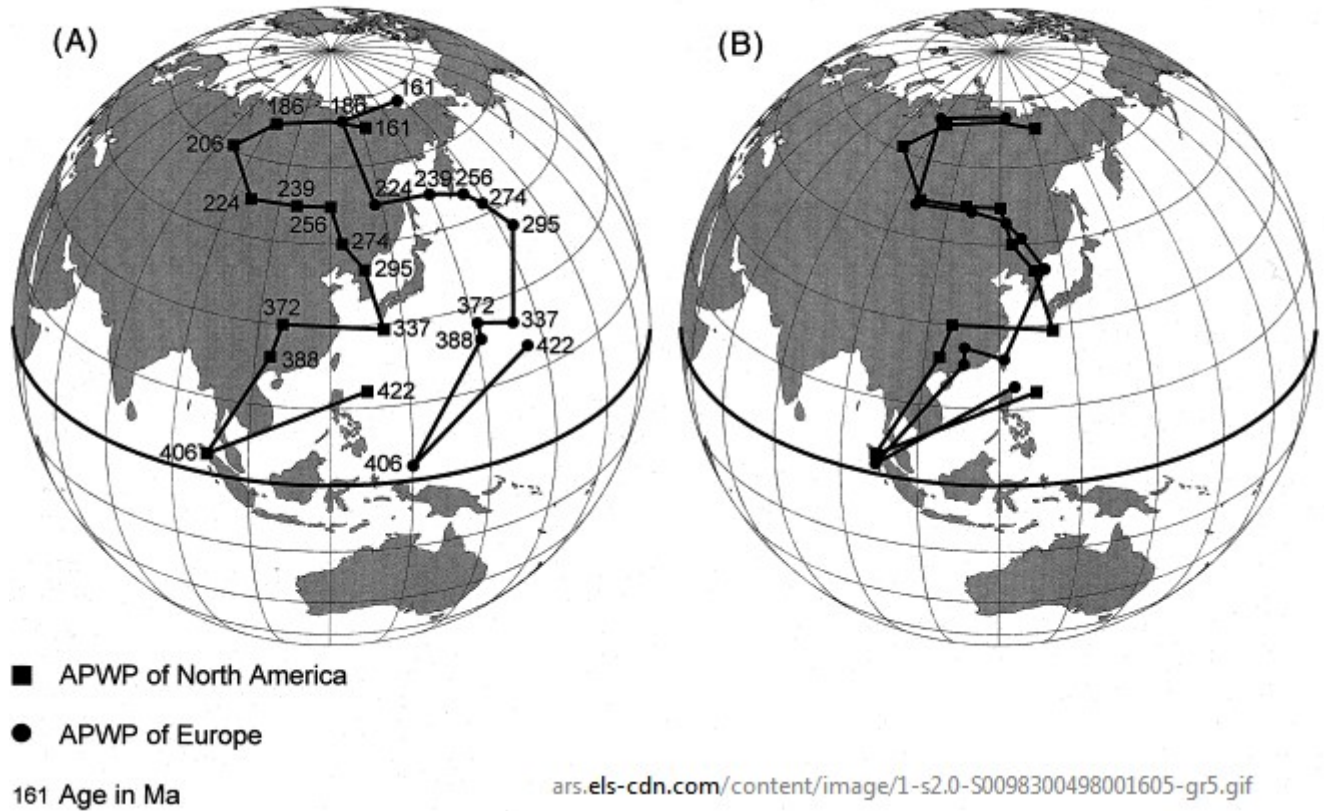


- By measuring the *inclination*, *declination*, and *intensity* of the paleomagnetic direction in stratum of different ages in different places, “polar-wandering paths” can be determined for different continents --- large, continuous land masses.

## Comparison of paleomagnetic data

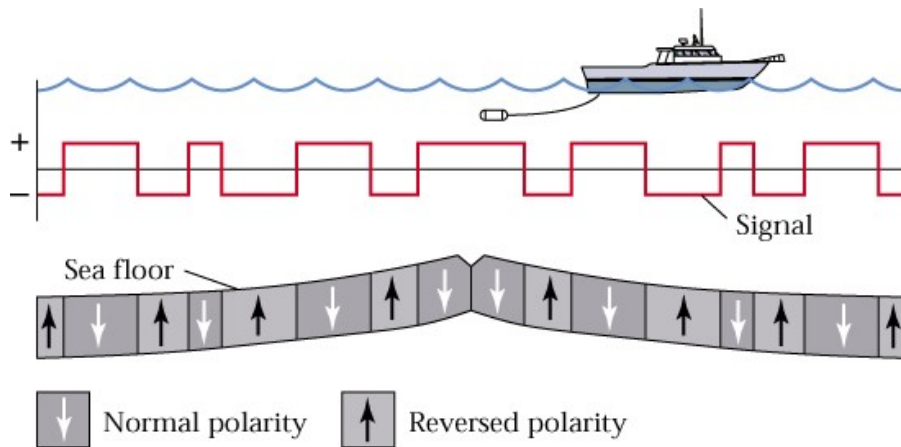
Different, but similar paths determined for different continents can be explained by:

- 1) ~~Having different poles for different continents~~
- 2) Magnetic poles have remained about the same while geographic poles and continents have moved

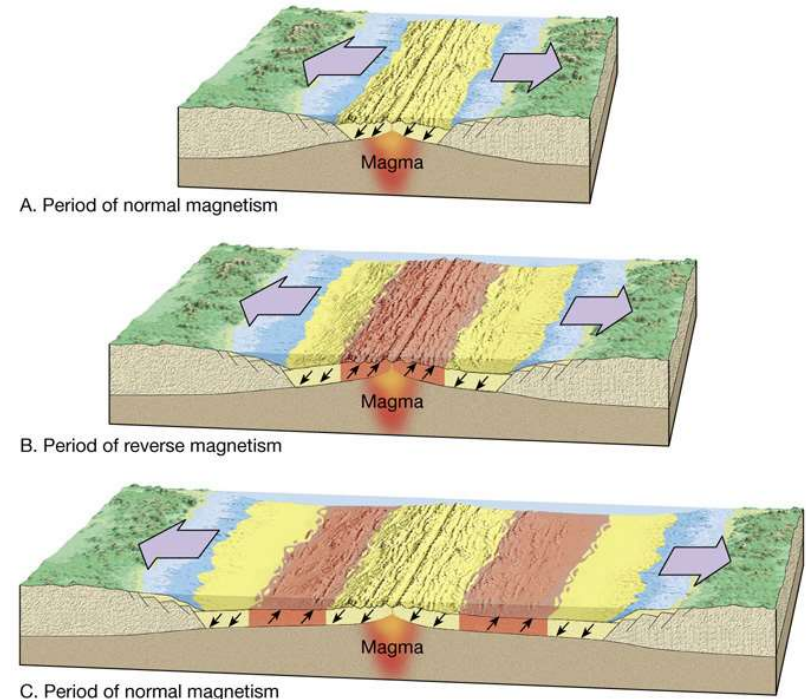


Magnetic surveys in the 1950's and 60's of oceanic crust revealed magnetic anomalies — reversals of magnetic polarity— and that these anomalies were disposed in symmetrical stripes paralleling the oceanic ridges.

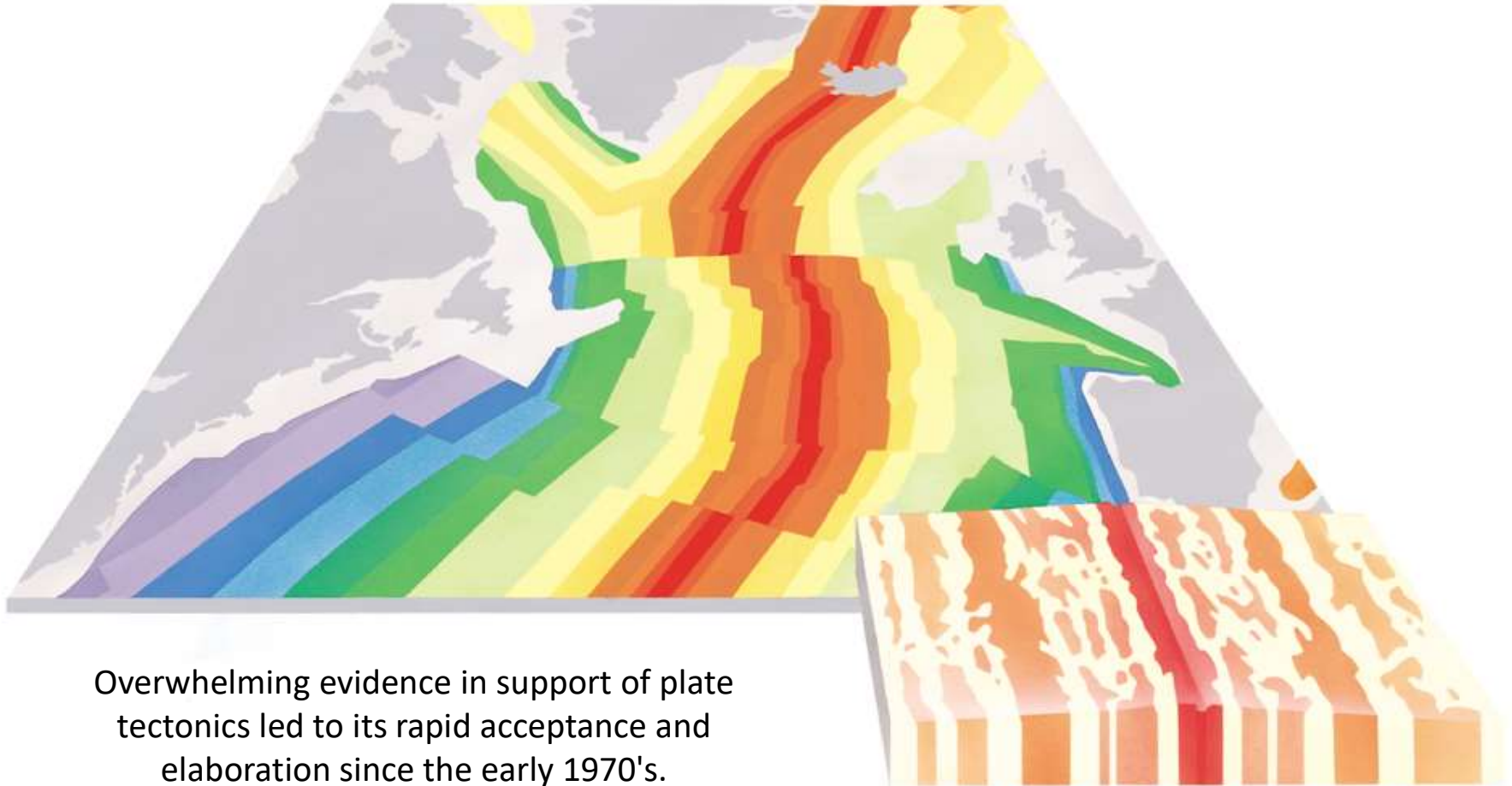
This indicates that new oceanic crust must be formed along the spreading ridges.



*This was a major breakthrough in plate-tectonic theory, because it shows that oceanic crust grows and can move with continental crust rather than having continents plowing through oceanic crust.*



Sea floor spreading is confirmed by the ages of fossils in sediments overlying oceanic crust of various ages, and radiometric dating of rocks on oceanic islands. These indicate that oceanic crust is youngest at the spreading ridges and oldest at the farthest points from the ridges.



Overwhelming evidence in support of plate tectonics led to its rapid acceptance and elaboration since the early 1970's.

***Note that oceanic crust is thinner than continental crust***

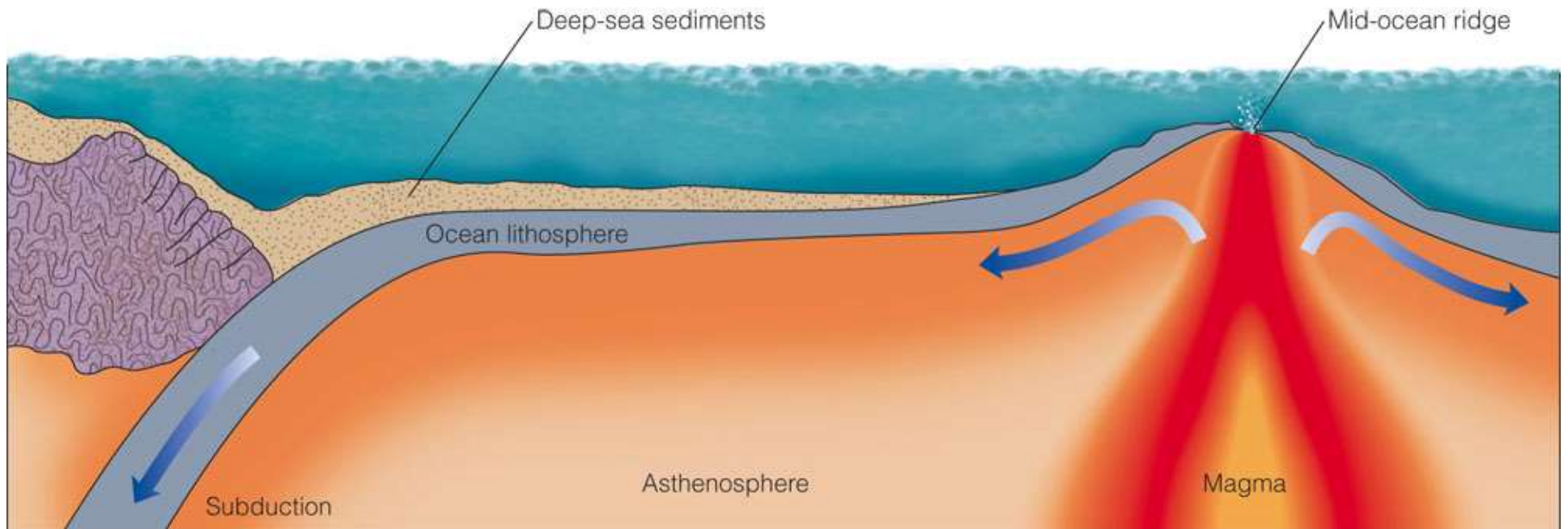
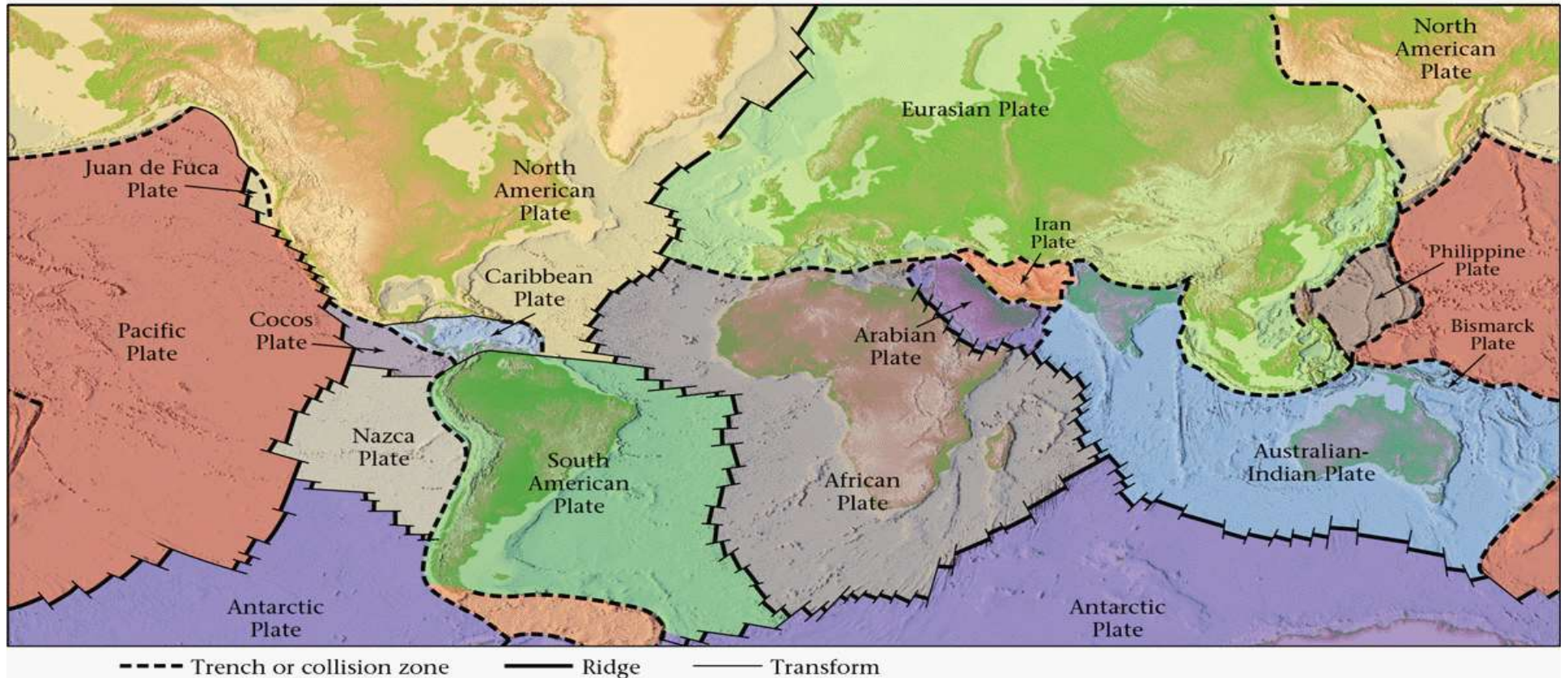


Plate-tectonic theory is widely accepted because it explains so many geologic phenomena, including volcanism, seismicity, mountain building, climatic changes, animal and plant distributions in the past and present, and the distributions of natural resources.

**For these reasons, it is known as a *unifying theory*.**

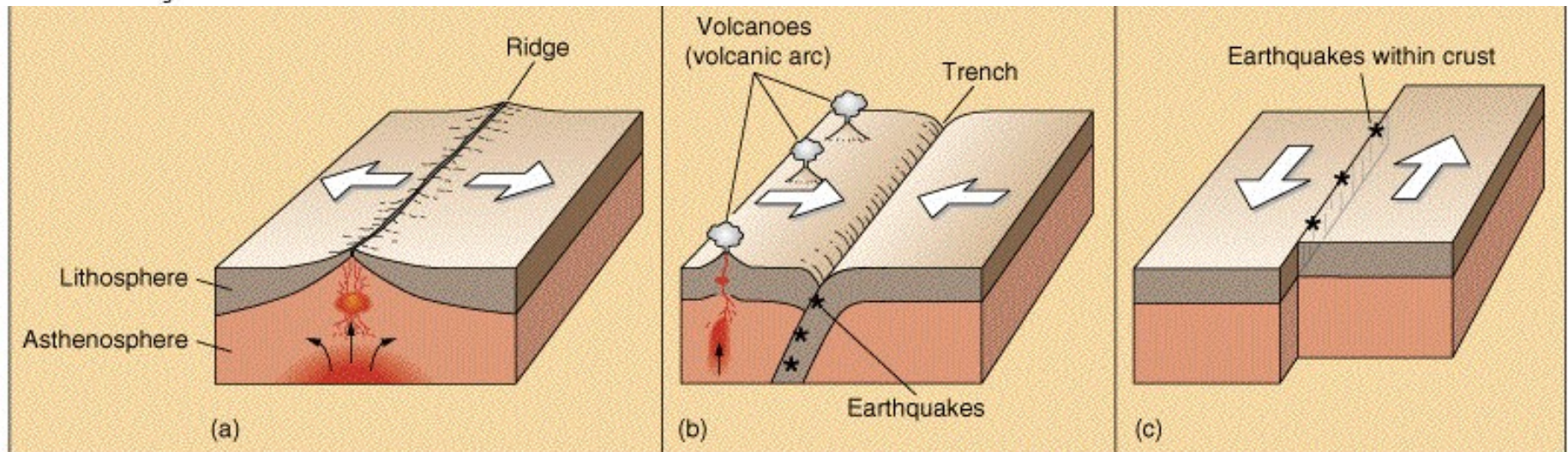


# There are three types of plate boundaries: divergent, convergent, and transform.

## *Types of Plate Boundaries*

Type	Example	Landforms	Volcanism
<b>Divergent</b>			
Oceanic	Mid-Atlantic Ridge	Mid-oceanic ridge with axial rift valley	Basalt
Continental	East African Rift Valley	Rift valley	Basalt and rhyolite, no andesite
<b>Convergent</b>			
Oceanic–oceanic	Aleutian Islands	Volcanic island arc, offshore oceanic trench	Andesite
Oceanic–continental	Andes	Offshore oceanic trench, volcanic mountain chain, mountain belt	Andesite
Continental–continental	Himalayas	Mountain belt	Minor
<b>Transform</b>	San Andreas fault	Fault valley	Minor

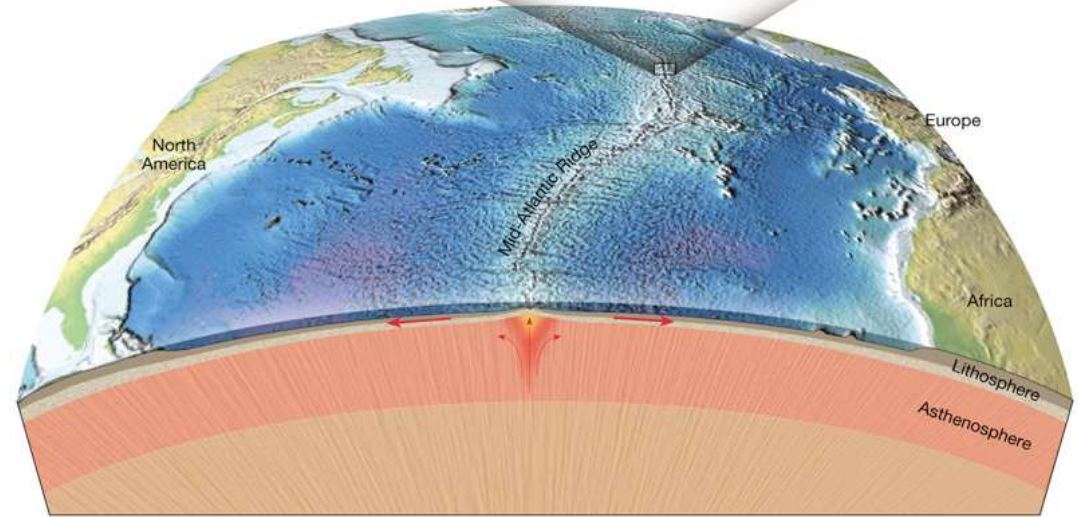
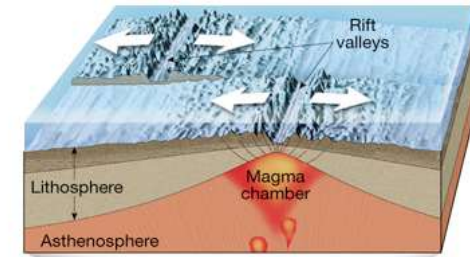
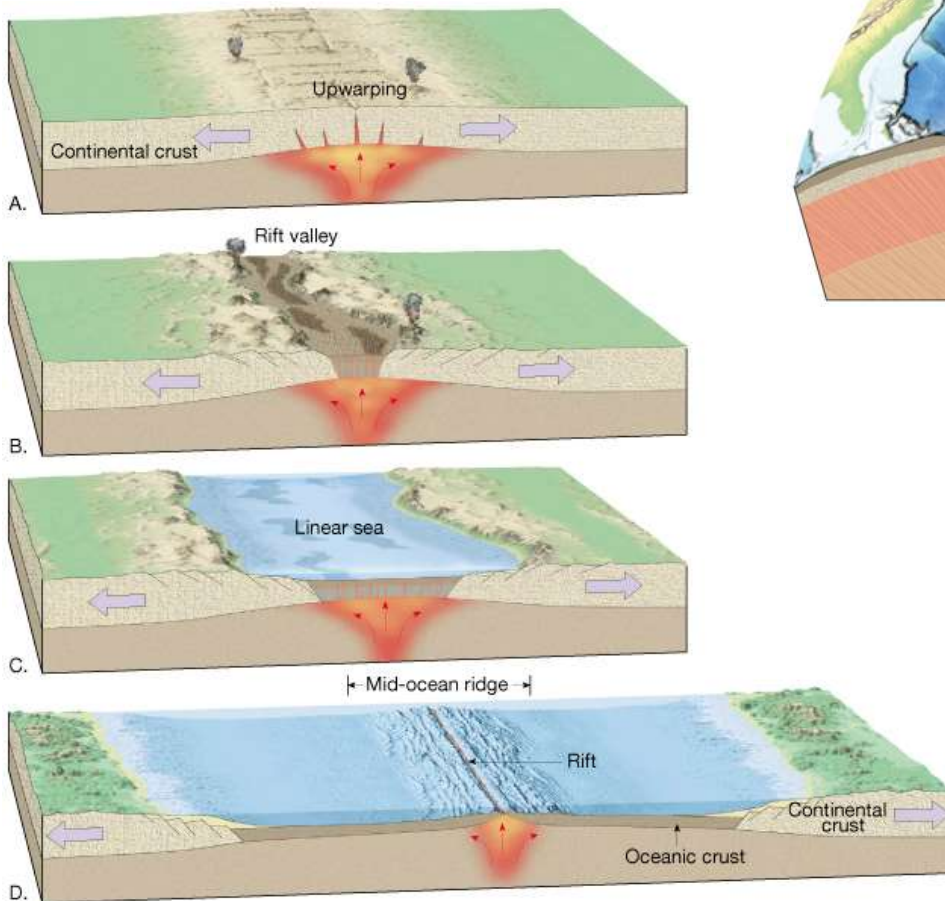
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[www.age-of-the-sage.org/tectonic\\_plates/volcanoes\\_earthquakes.gif](http://www.age-of-the-sage.org/tectonic_plates/volcanoes_earthquakes.gif)

**Divergent boundaries are where plates move away from each other and new crust is formed along ocean ridges and in continental basins**

Type 2 continental

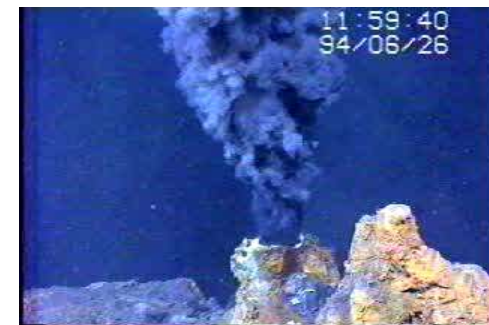
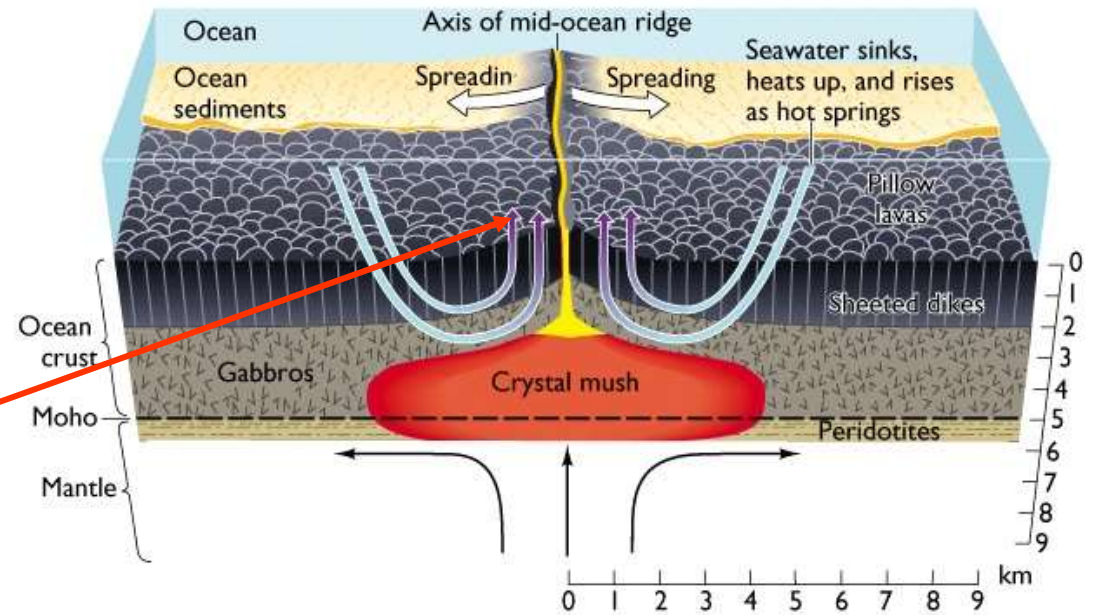
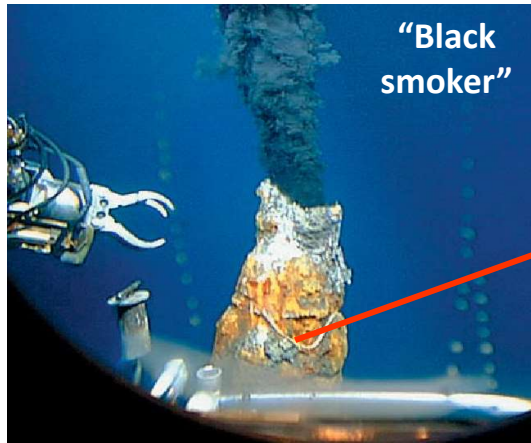


Type 1 oceanic

Two different types but have something in common

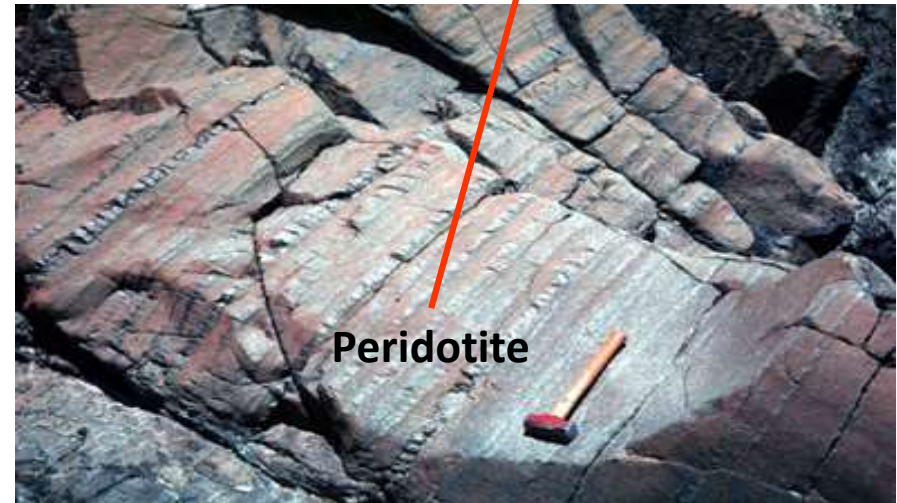
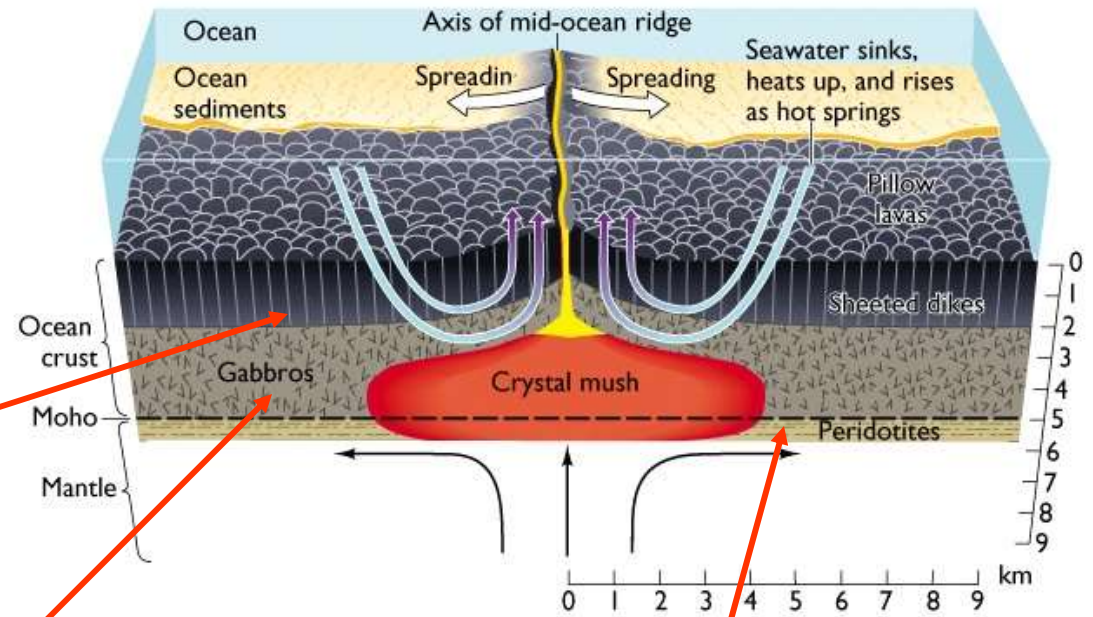


# Oceanic divergent boundaries



Precambrian basalt from 2-billion-year-old slice of seafloor, Quebec, Canada

# Oceanic divergent boundaries

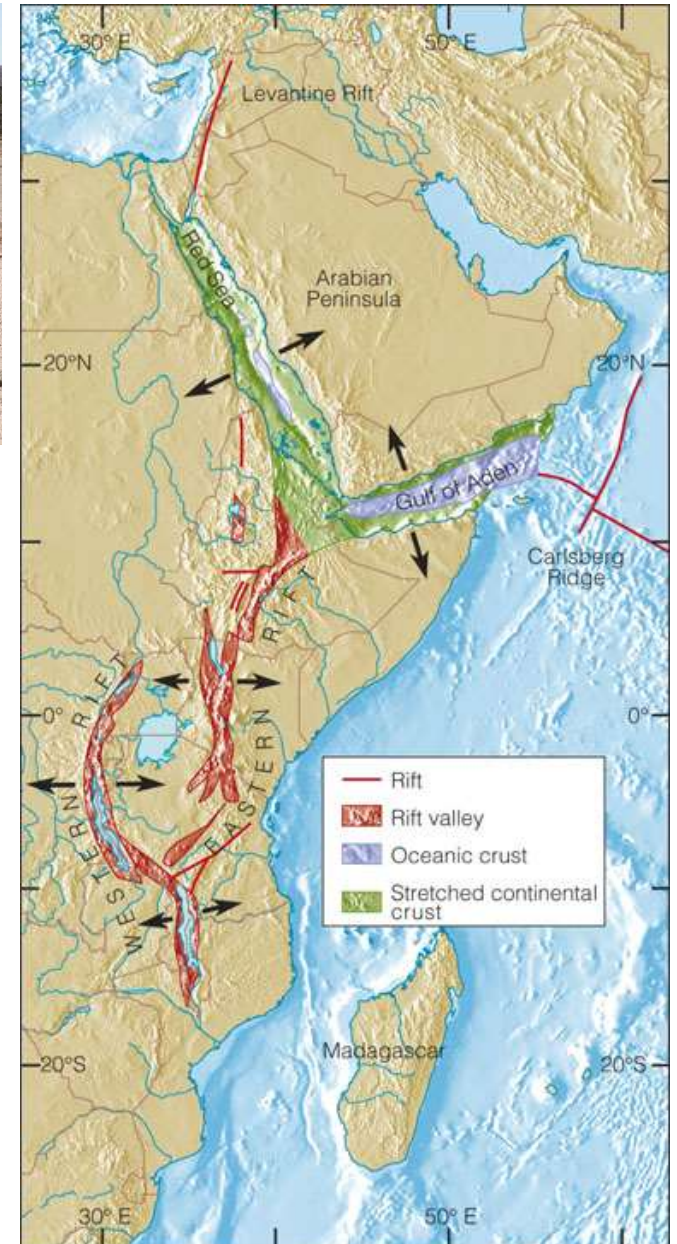


# Continental divergent boundaries

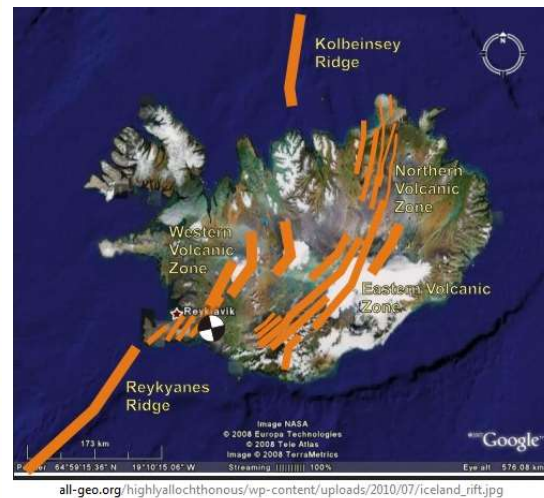
Wright, T.J., C. Ebinger, J. Biggs, A. Ayele, G. Yirgu, D. Keir, A. Stork, 2006, Magma-maintained rift segmentation at continental rapture in the 2005 Afar dyking episode, Nature, 442, 291-294



# East-African rift



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# Icelandic rifts

# Continental divergent boundaries – closer to home

Eastern North American continental rift system and the Newark basin

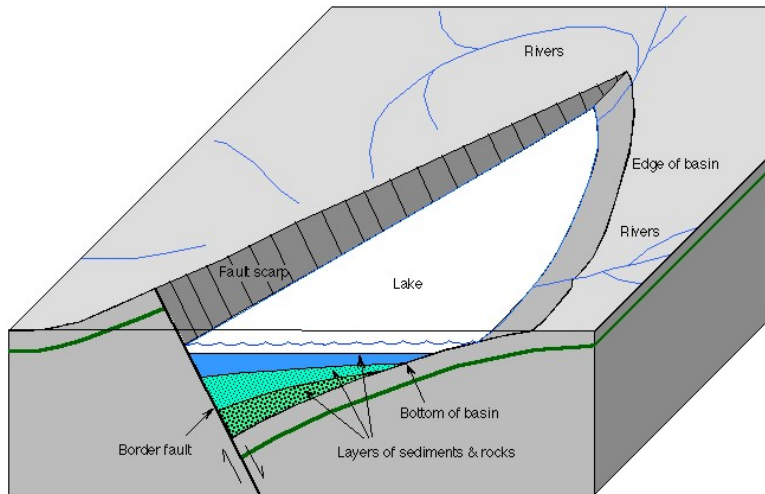
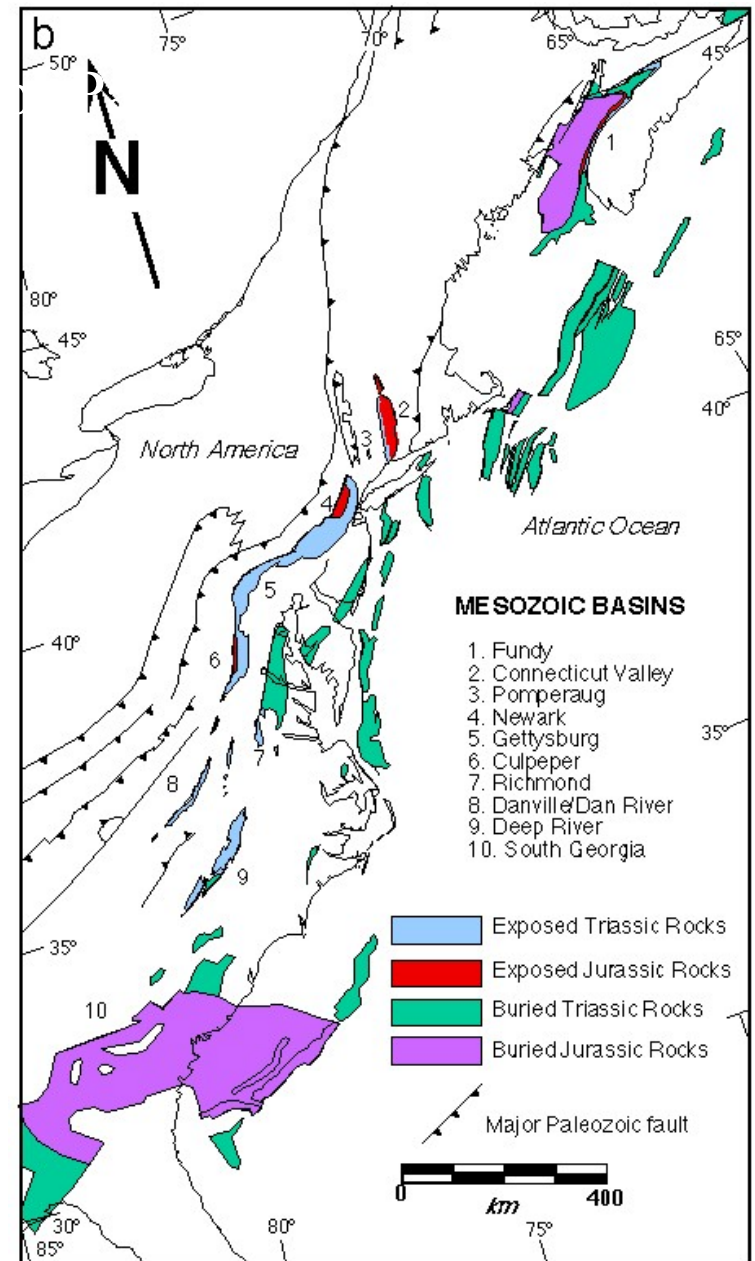
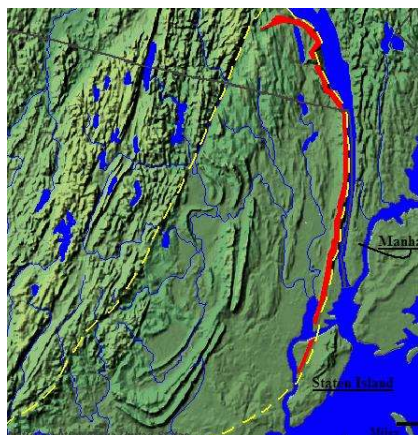
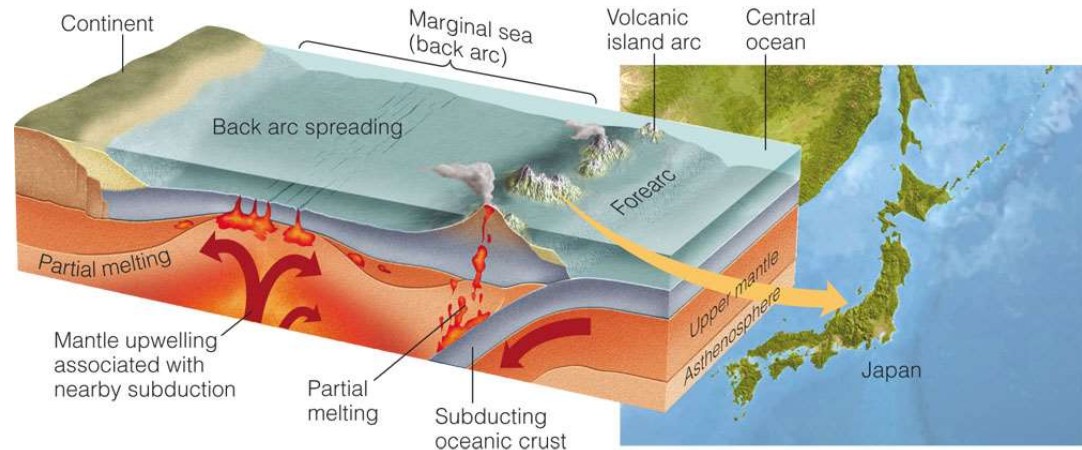


Figure 4: Cutaway block diagram of a rift basin. Note the half-graben geometry (triangular) in the cross section view (front panel).

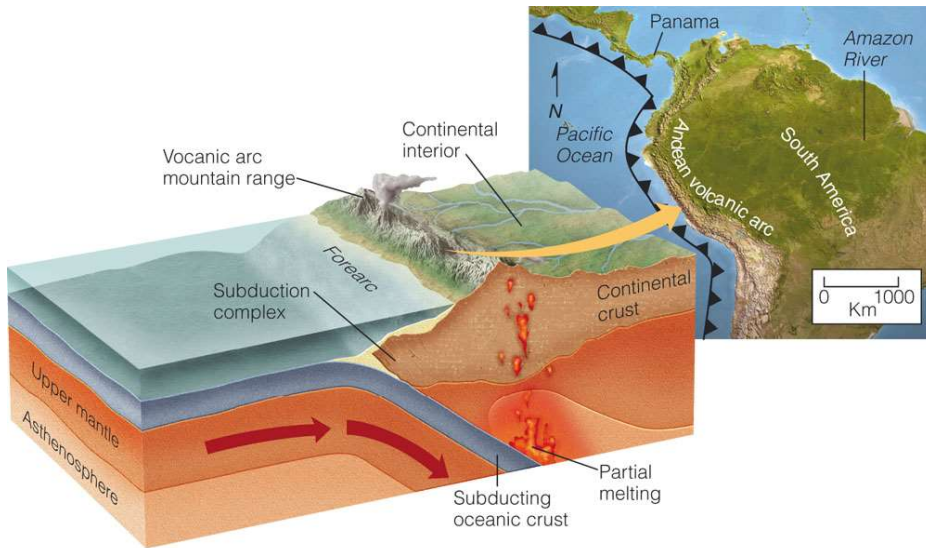


**Convergent boundaries exist where one plate is *subducted* beneath another and crust is destroyed, and/or two plates with continental crust on their leading edges collide, and mountains are formed**



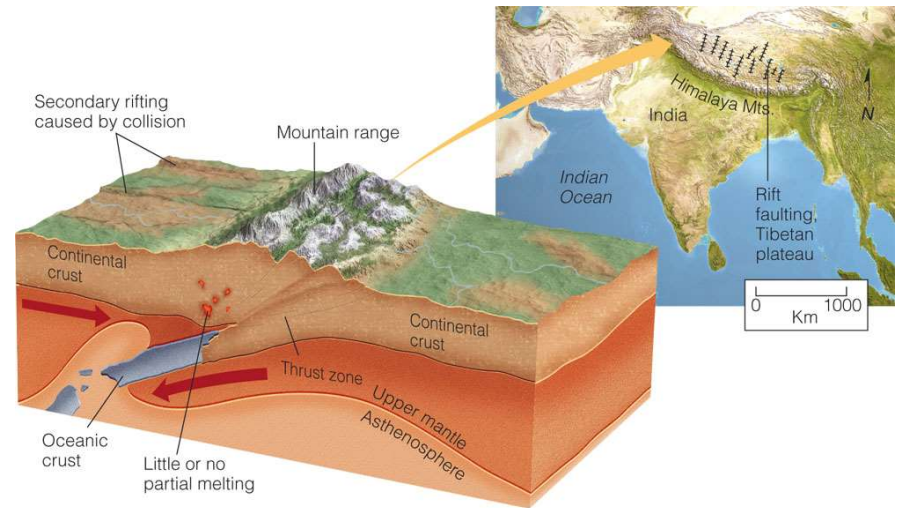
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**Type 1 ocean – ocean**



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**Type 2 ocean – continent**



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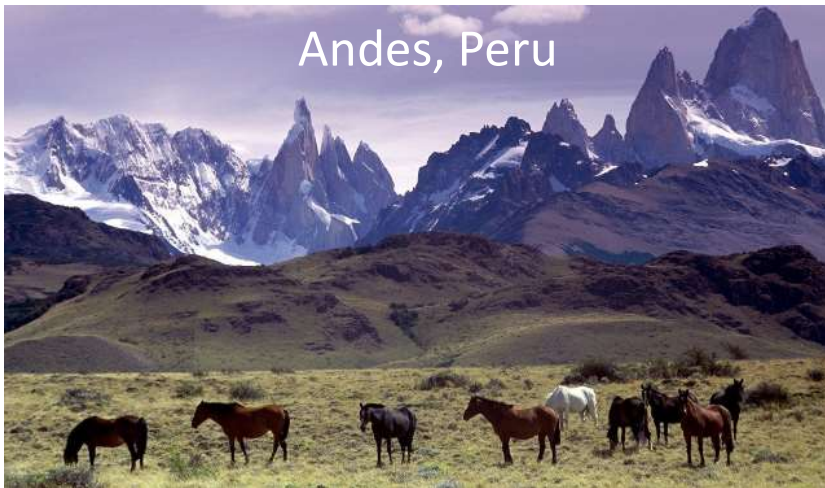
**Type 3 continent – continent**

## Japanese Alps



[www.info-toyama.com/english/tlibrary/img/nature/toyama/3\\_003.jpg](http://www.info-toyama.com/english/tlibrary/img/nature/toyama/3_003.jpg)

## Andes, Peru



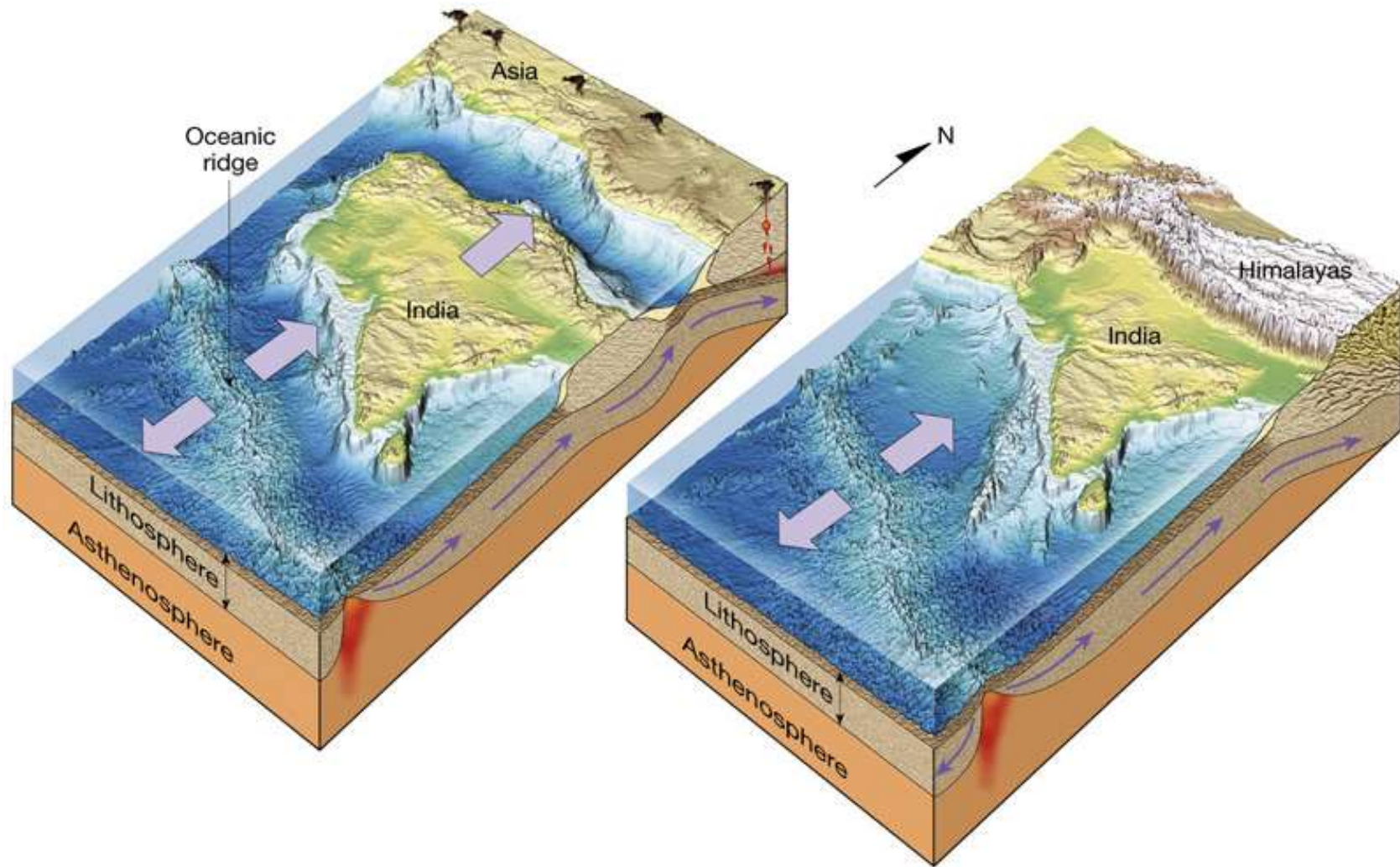
[www.openwalls.com/image/19113/horses\\_andes\\_mountains\\_argentina\\_\\_1400x1050.jpg](http://www.openwalls.com/image/19113/horses_andes_mountains_argentina__1400x1050.jpg)

## Himalayas, Mongolia - China



[www.socialmediastrategiessummit.com/blog/wp-content/uploads/Mount\\_Everest.jpg](http://www.socialmediastrategiessummit.com/blog/wp-content/uploads/Mount_Everest.jpg)

# Himalayas, Mongolia - China

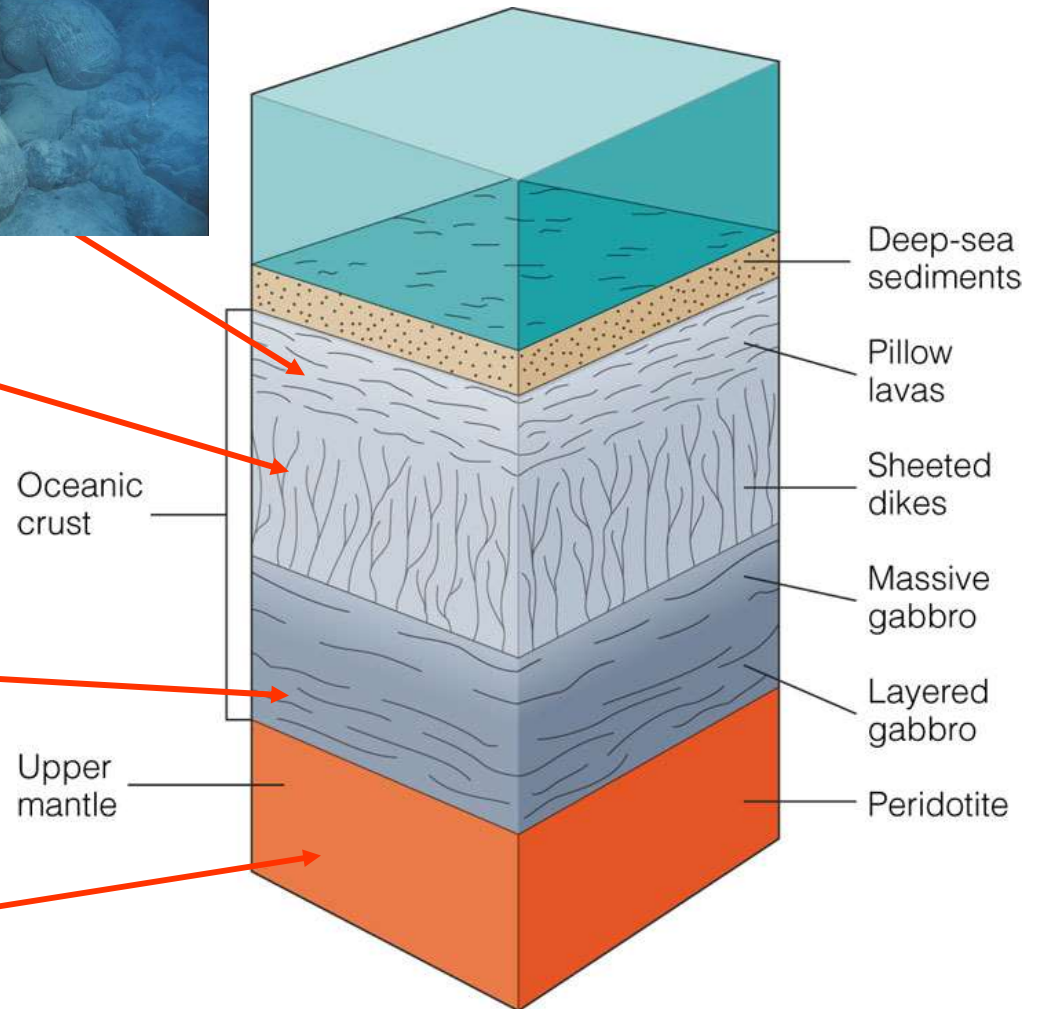




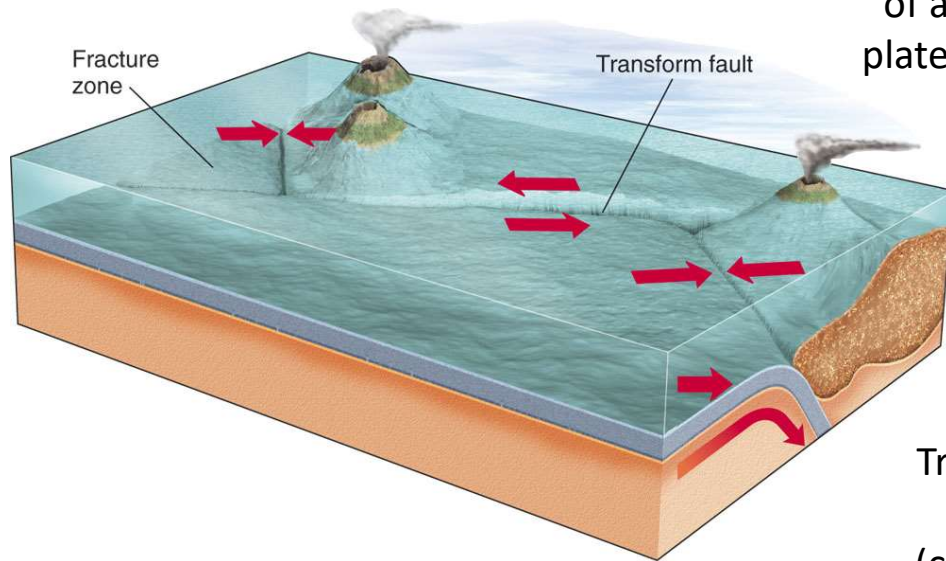
**Examples of alpine folding and faulting in continental mountain belts**



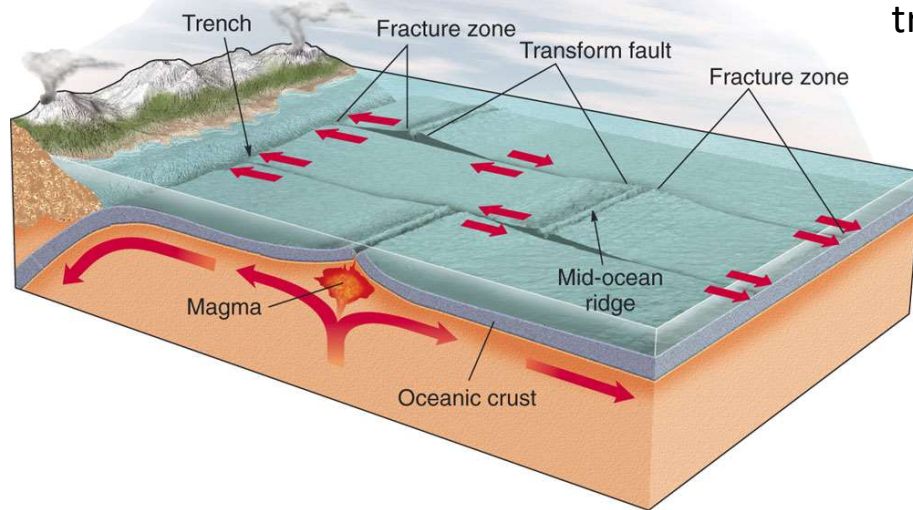
Ancient, oceanic crustal rocks that formed at divergent boundaries that are locally thrust onto land (obducted) at oceanic-continental convergent plate boundaries, are called ophiolites.



**Transform boundaries are where the margins of two plates slide horizontally past one another.**



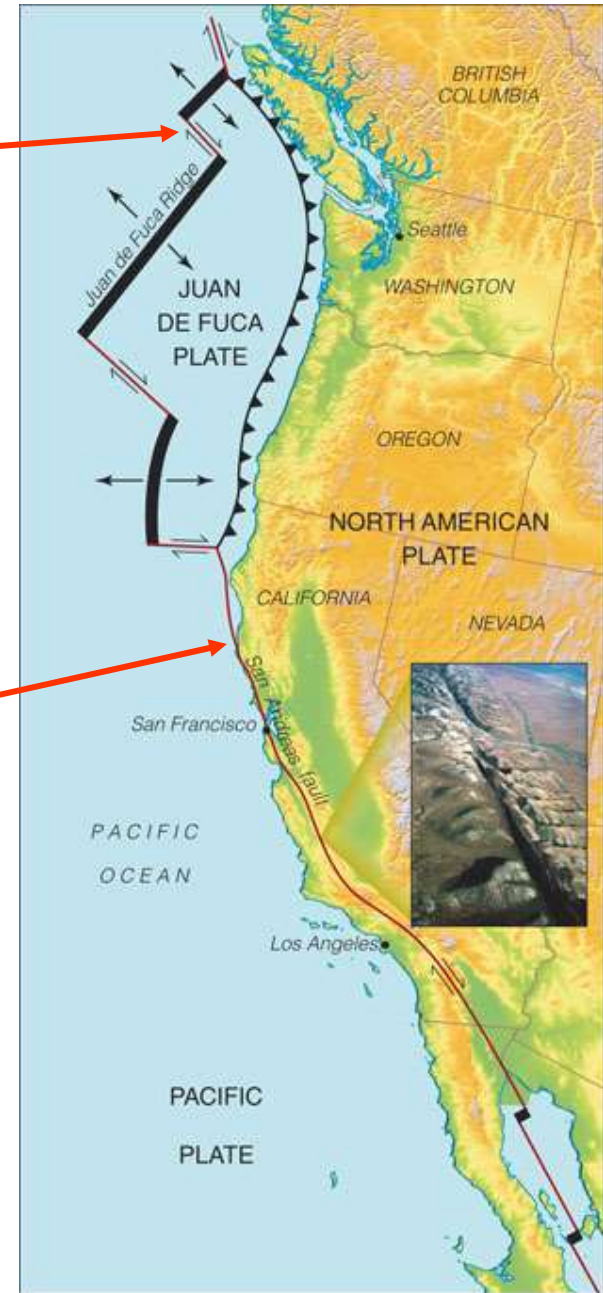
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Transform faults (oceanic) as part of a complex plate boundary

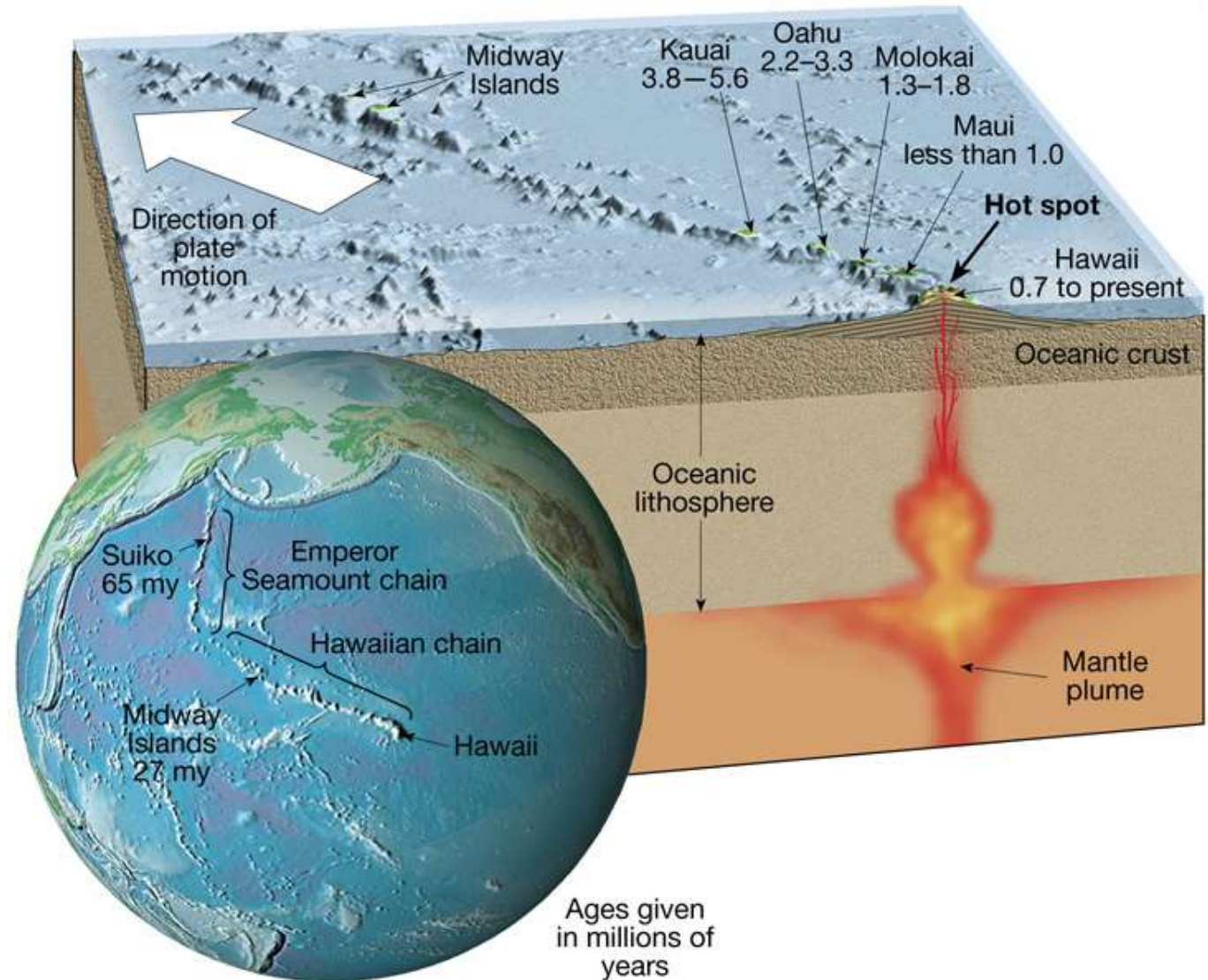
Transcurrent fault (continental) transform plate boundary



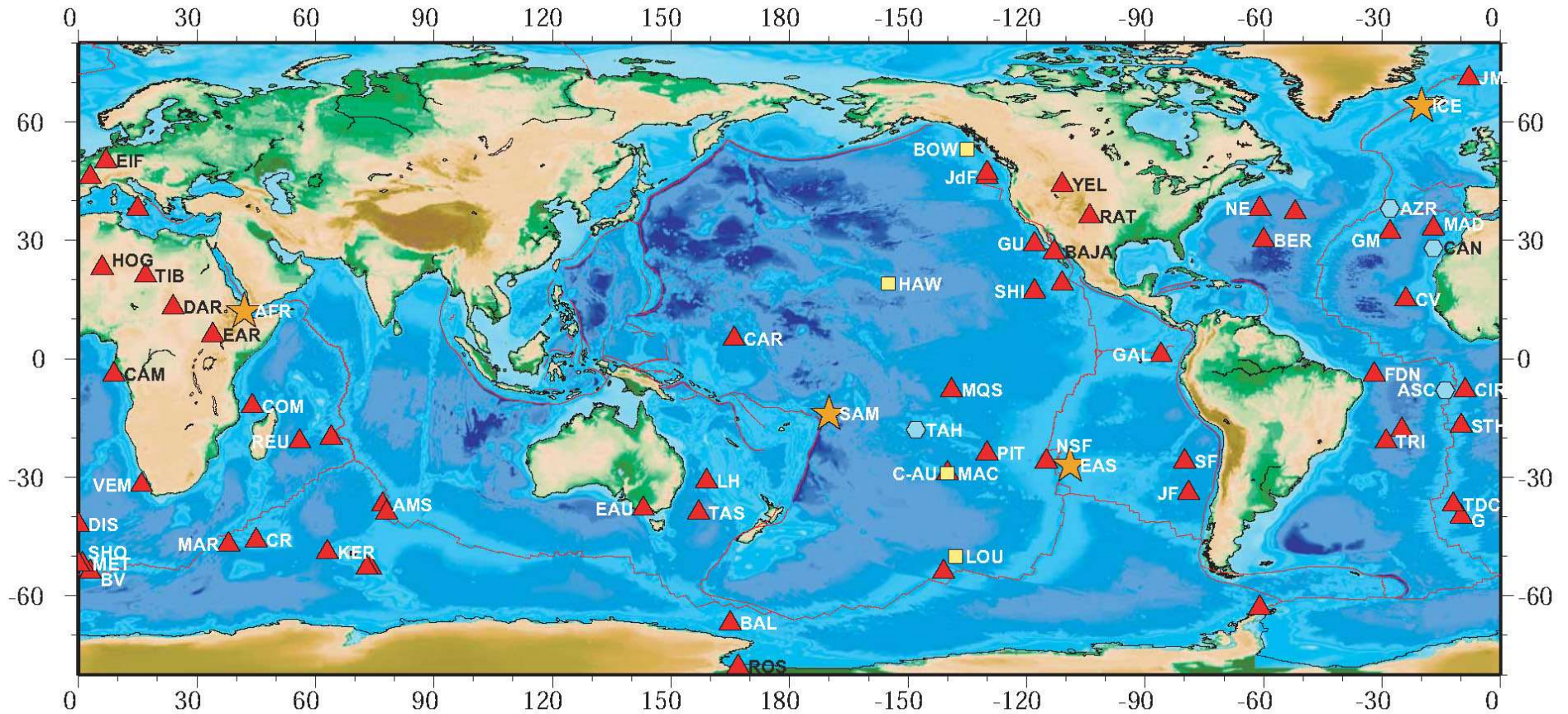
© 2007 Thomson Higher Education

# Hot Spots – An intraplate feature

A location on Earth's surface where a stationary (?) column of magma, originating deep within the mantle (mantle plume), has risen to the surface to form a volcano.



# Hot Spots – An intraplate feature

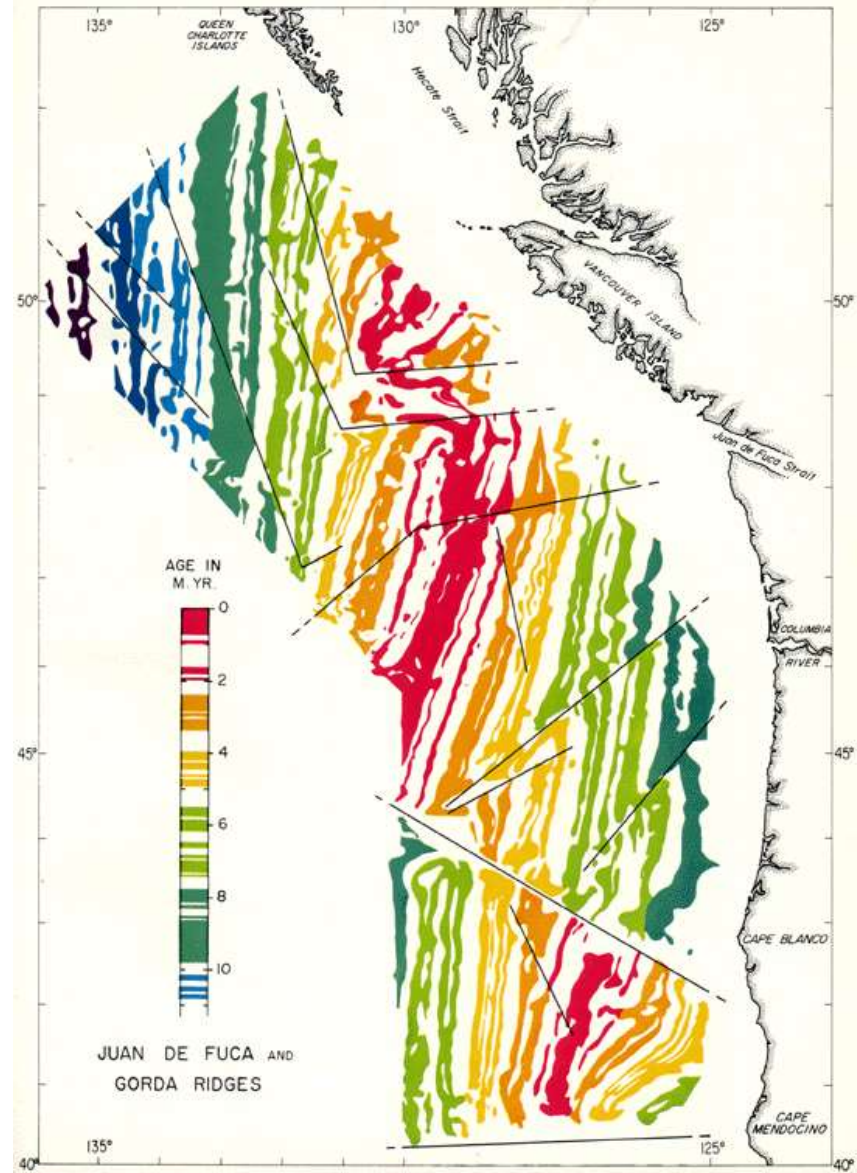


[www.faculty.gg.uwo.edu/ducker/Yellowstone/YEL%20PLUME%20TALK%202005\\_files/slide0036\\_image003.jpg](http://www.faculty.gg.uwo.edu/ducker/Yellowstone/YEL%20PLUME%20TALK%202005_files/slide0036_image003.jpg)

# How do we measure plate motions?

Old ways:

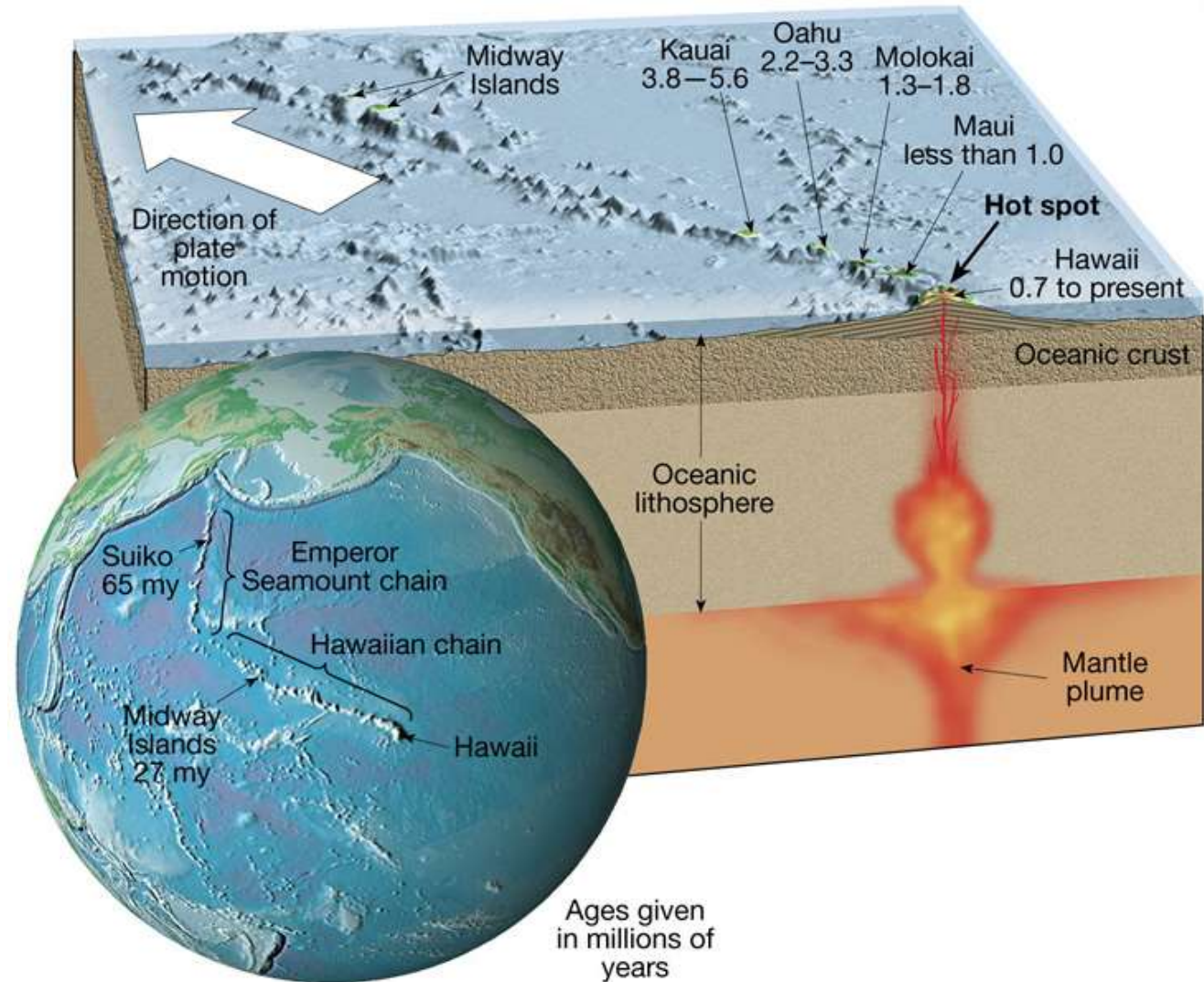
- 1) Measure sediment thickness on ocean plates and divide by distance from a ridge (least accurate)
- 2) Dating magnetic anomalies on the ocean floor. Knowing the age of the anomaly, and the distance between them, divide the distance by the difference of the bounding ages.
- 3) Similar to No. 2 but using chains of hotspots.



# How do we measure plate motions?

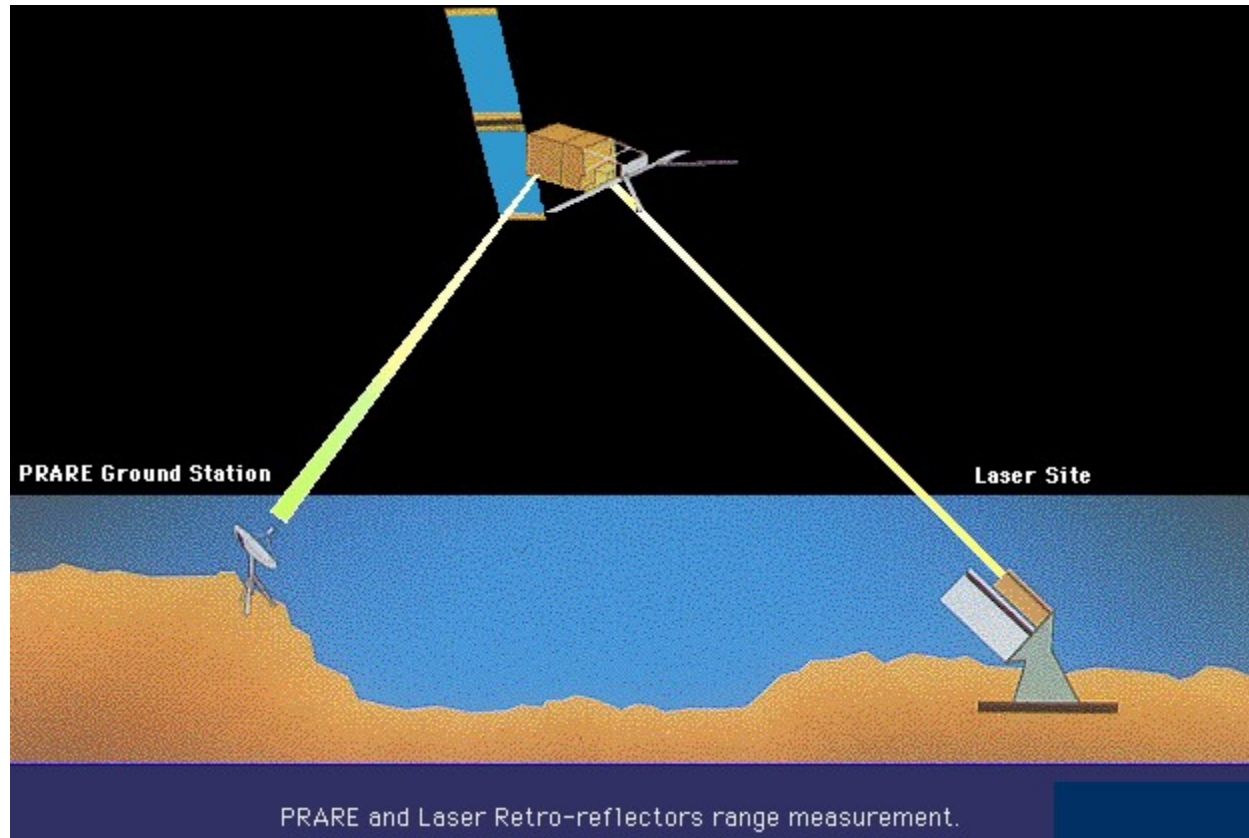
Old ways:

- 3) Similar to No. 2 but using oceanic-mountain chains formed by a hotspot



# How do we measure plate motions?

New ways: 1) Laser – satellite ranging



[www.earth.esa.int/rootcollection/eeo4.10075/\\_PRARE\\_and\\_Laser\\_RETRO-Reflectors\\_range\\_measurements\\_Image.gif](http://www.earth.esa.int/rootcollection/eeo4.10075/_PRARE_and_Laser_RETRO-Reflectors_range_measurements_Image.gif)

# How do we measure plate motions?

New ways:

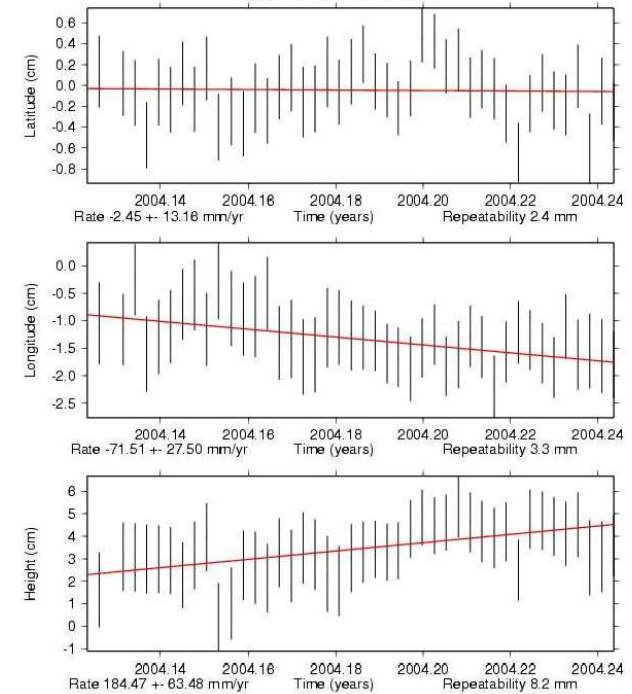
2) Ground-based, Global-Positioning-Systems (GPS)



[http://www.nist.gov/pml/div688/grp40/images/GPS\\_Constellation\\_2.gif](http://www.nist.gov/pml/div688/grp40/images/GPS_Constellation_2.gif)



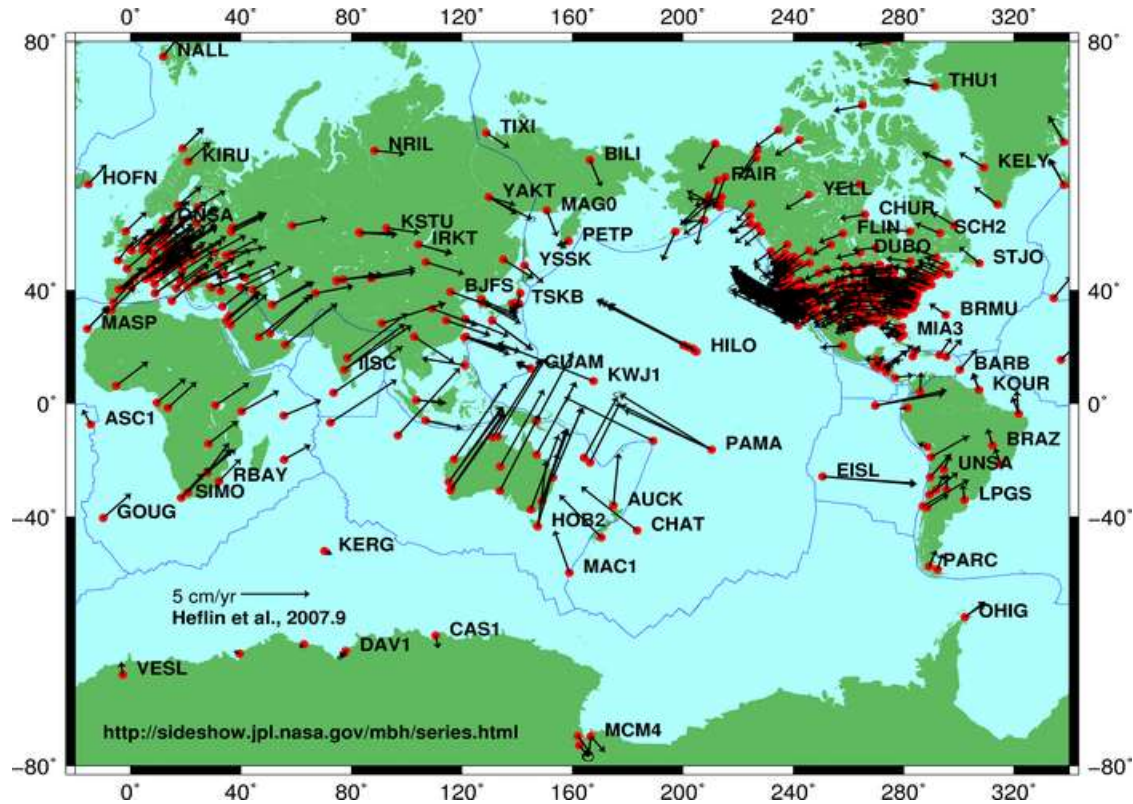
Time series for MANZ.



<http://sideshow.jpl.nasa.gov/mbh/series.html>

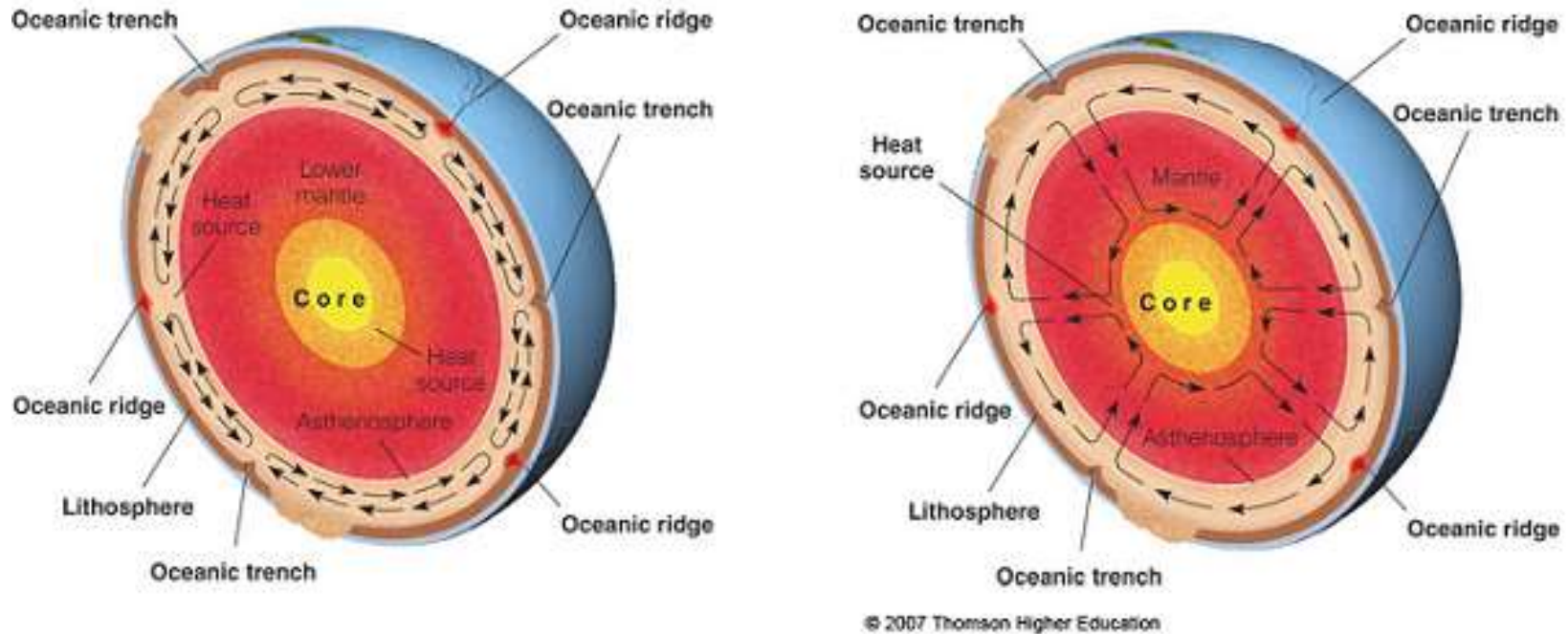


# How fast are plate velocities?



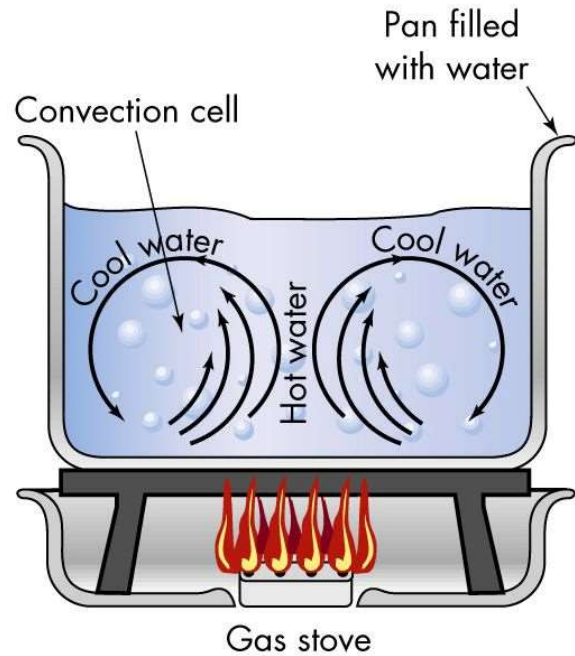
For 909 stations, 1 mm to 79 mm per year horizontal velocity, average 26 mm/yr

# What drives plate motions?

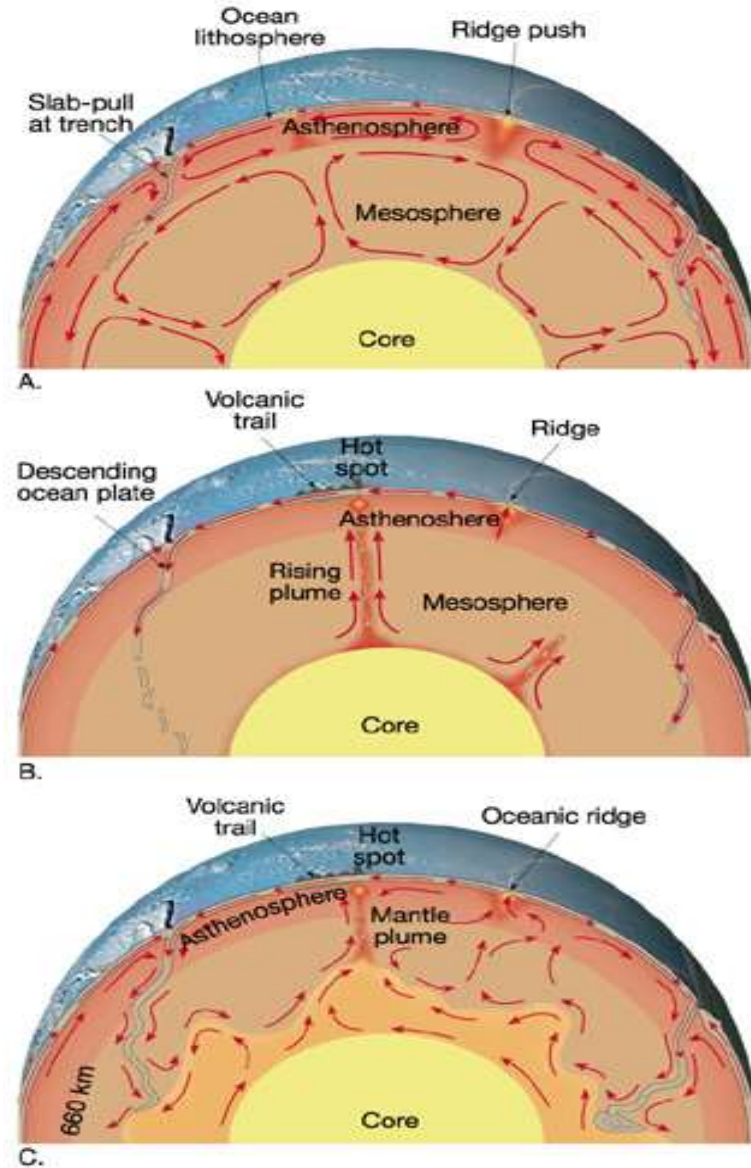


- 1) Thermal convection cells act as a 'conveyor-belt, carrying the crust along

# What drives plate motions?

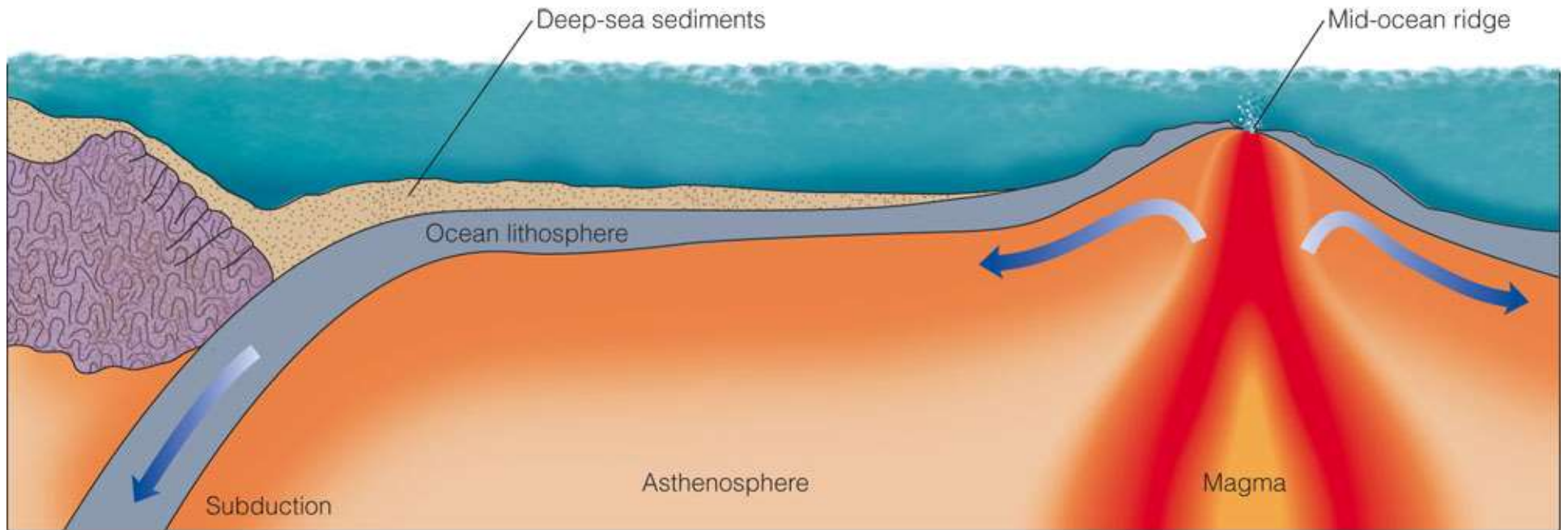


Various models for mantle convection have been proposed



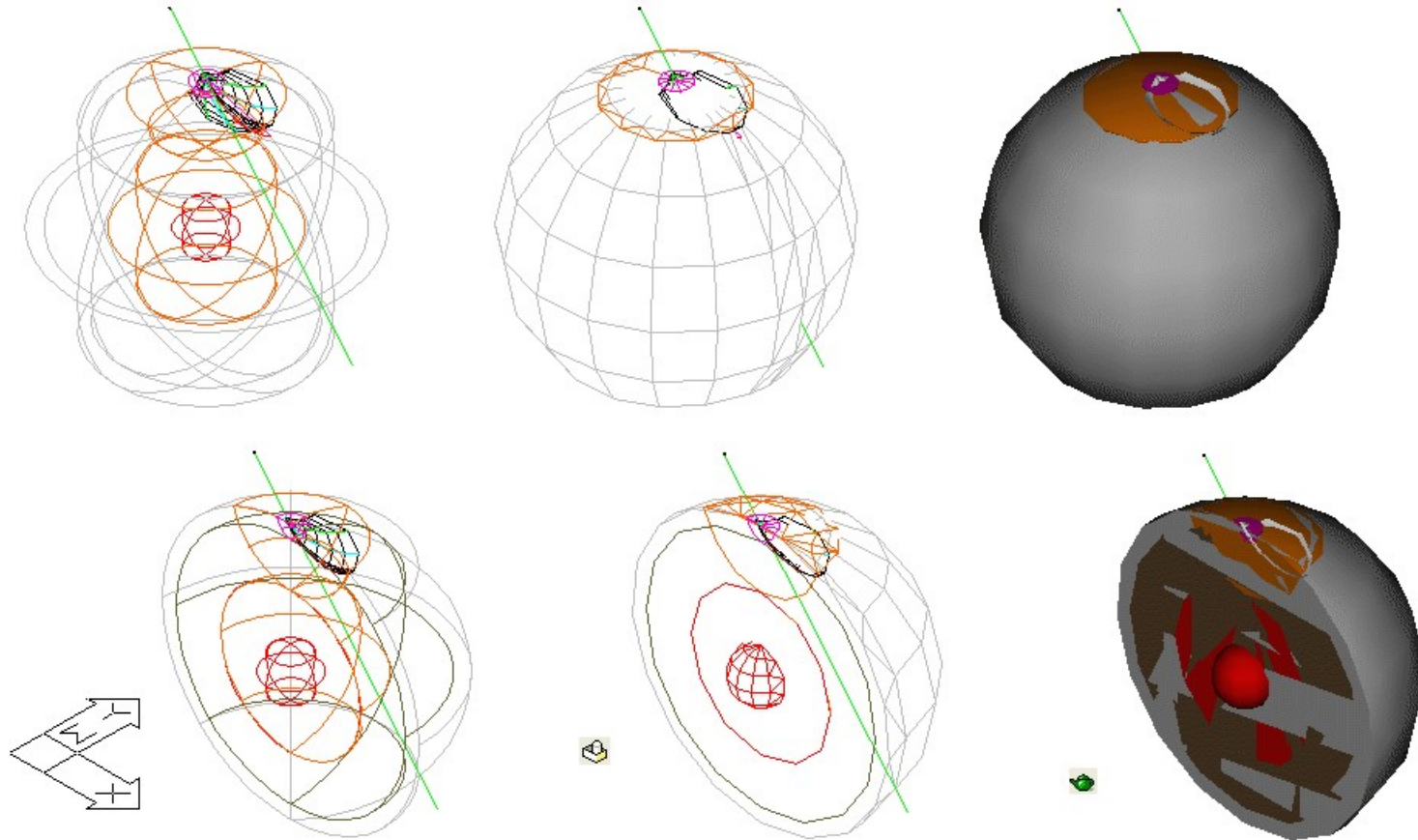
# What drives plate motions?

- 2) Gravity including the effects of 'ridge push' and 'slab pull'



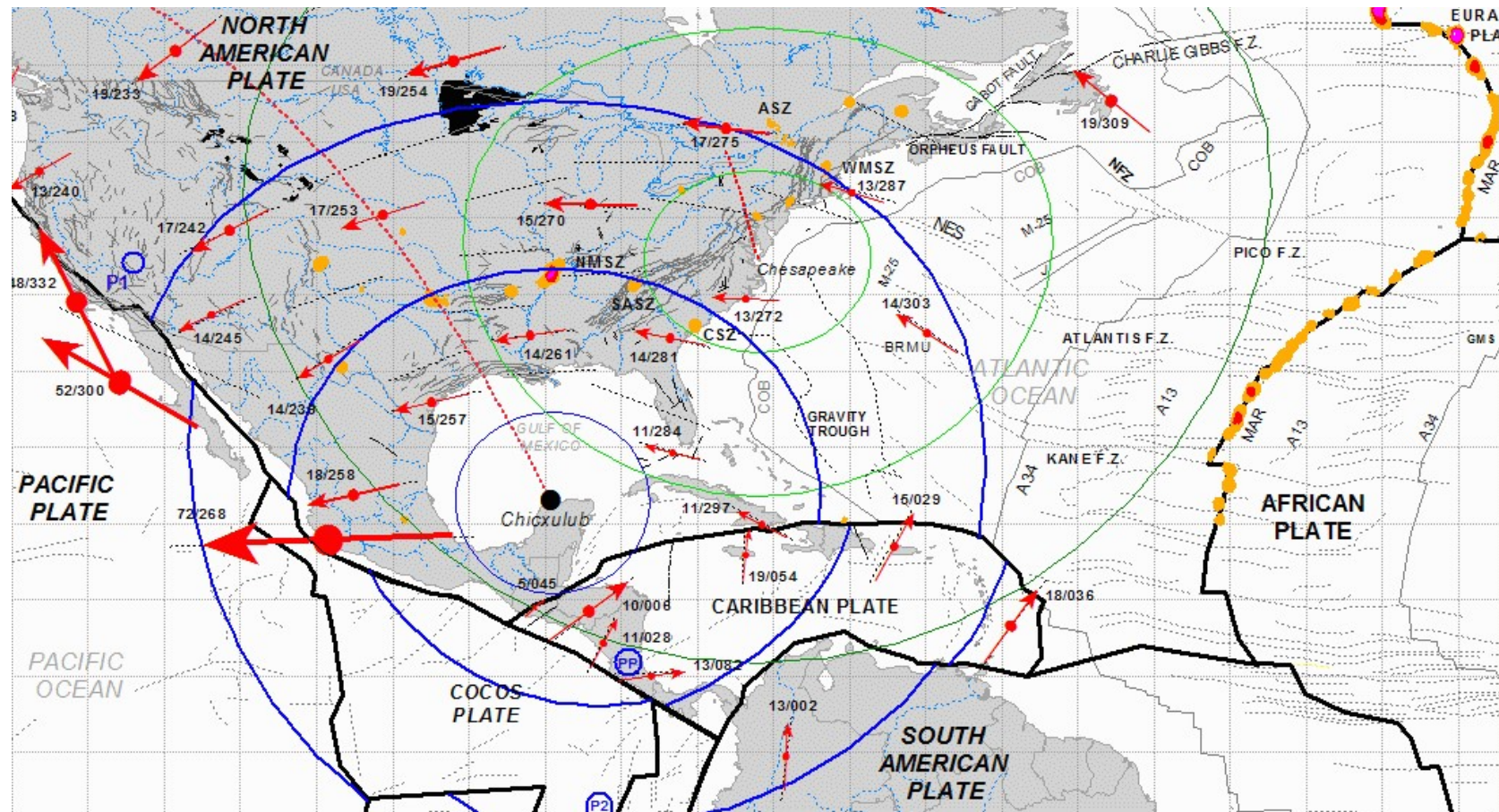
# What drives plate motions?

3) Directed push from, oblique, hypervelocity (> 3km/sec to 30 km/sec), bolide impacts ?



# What drives plate motions?

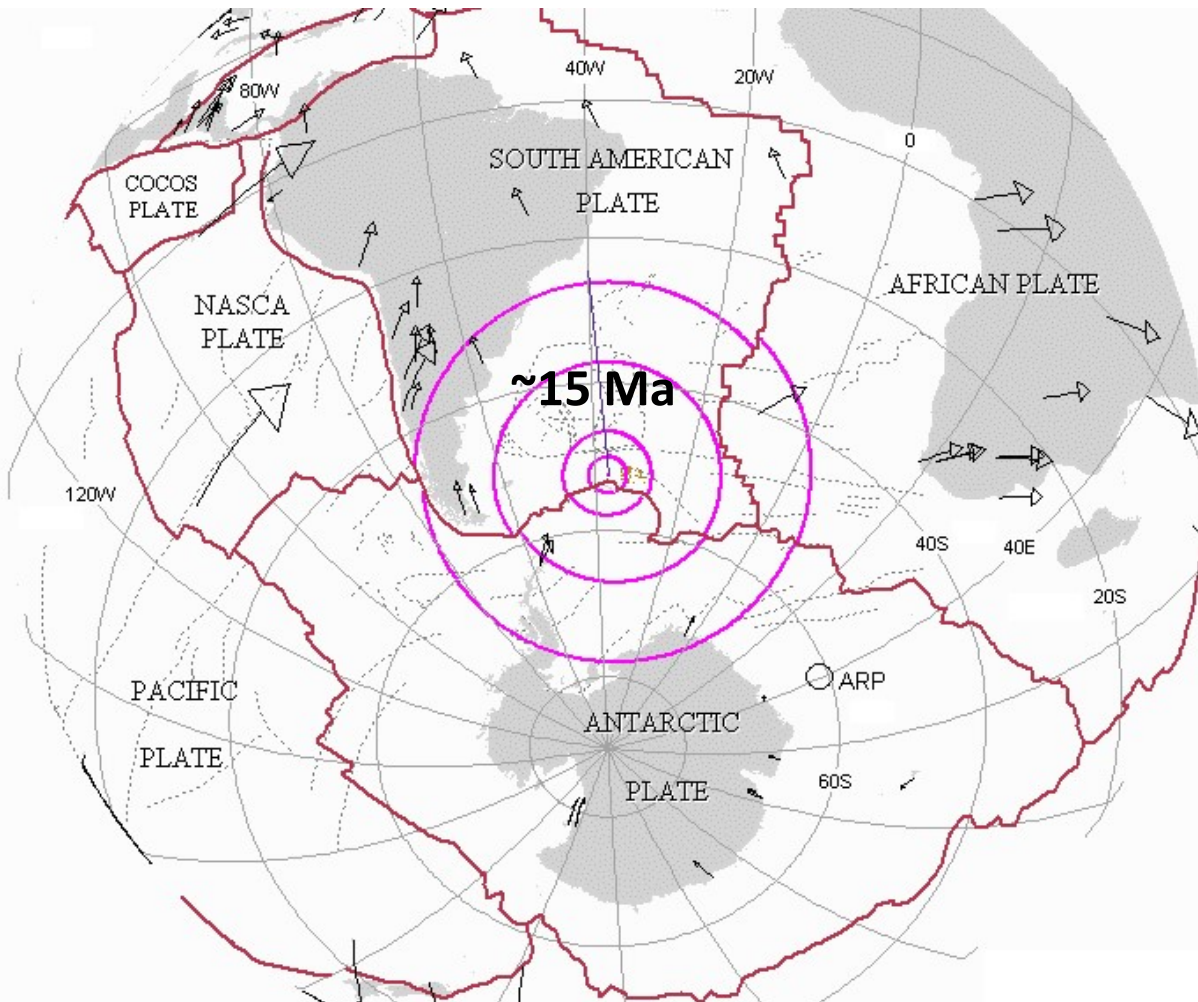
3) Directed push from, oblique, hypervelocity ( $> 3\text{ km/sec}$  to  $30\text{ km/sec}$ ), bolide impacts ?



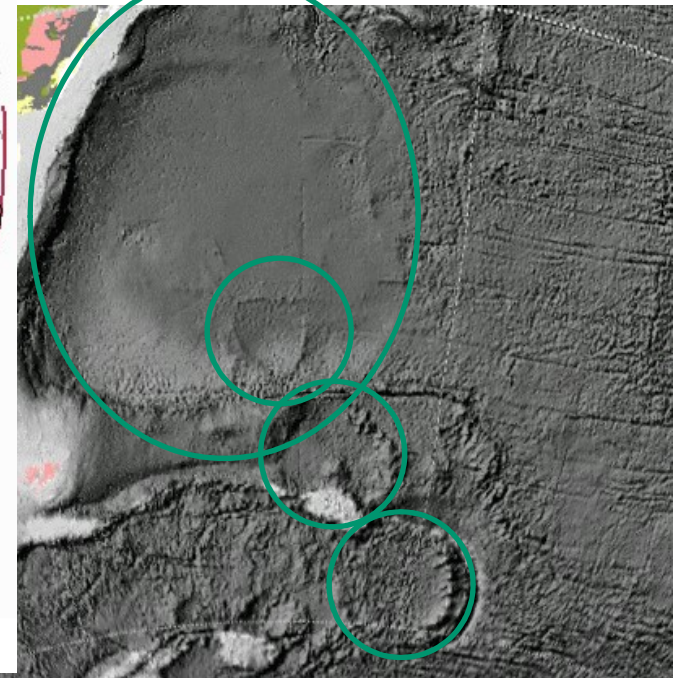
*The North American and Caribbean plates move in concert around the Chicxulub impact*

# What drives plate motions?

3) Large, hypervelocity (> 3km/sec to 30 km/sec) bolide impacts ?

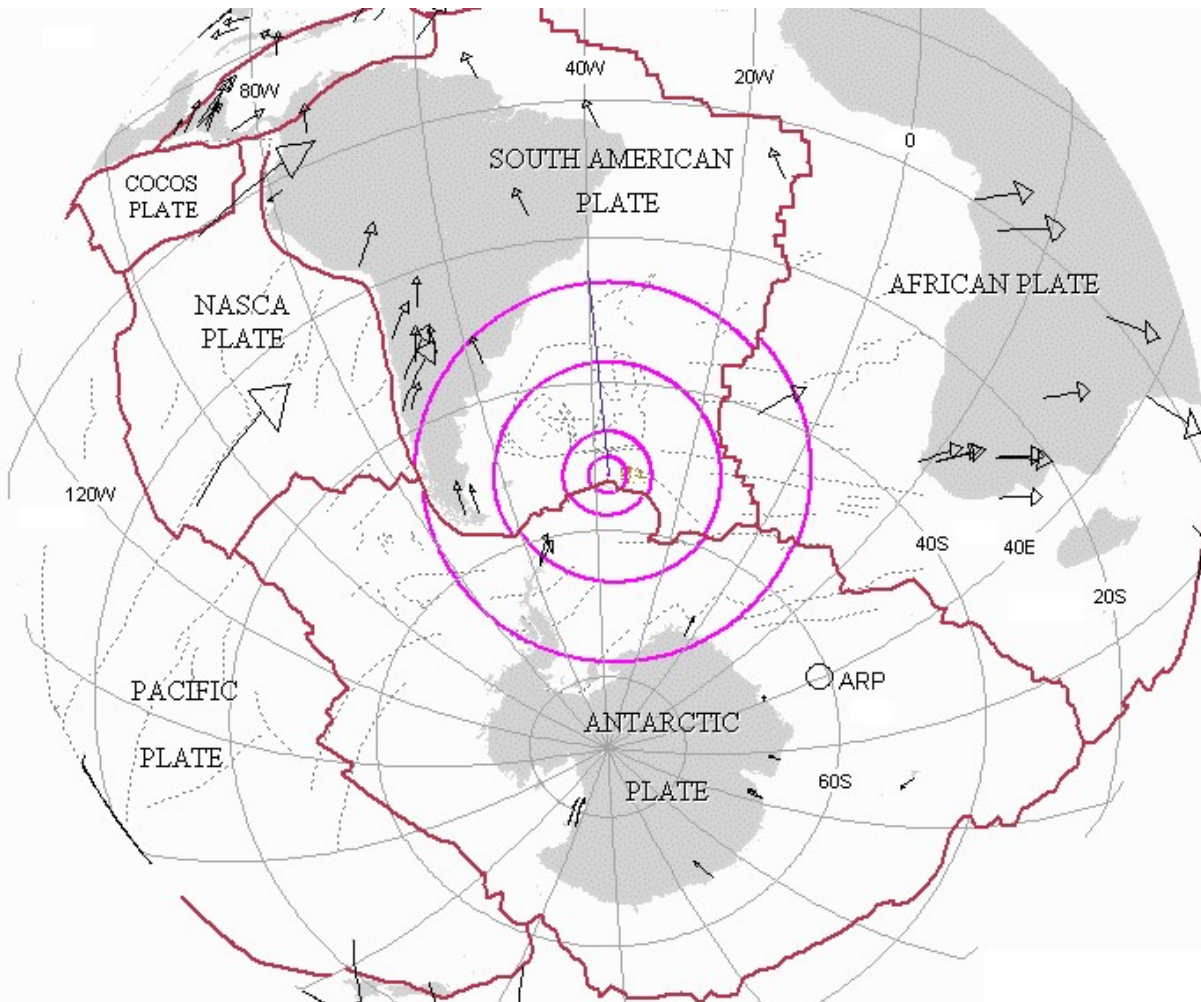


The South American Plate is moving northward away from a suspected, ~15 Ma impact crater in the South Atlantic

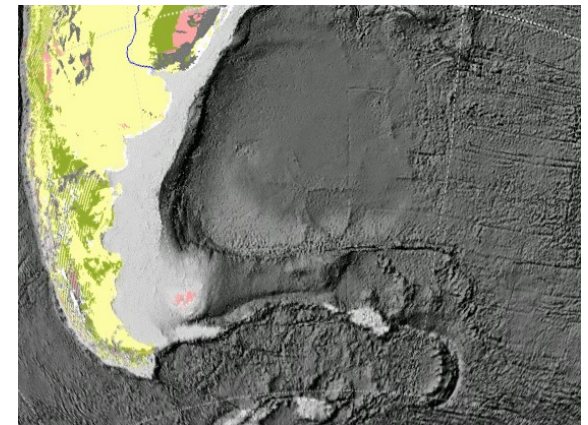


# What drives plate motions?

3) Large, hypervelocity (> 3km/sec to 30 km/sec) bolide impacts ?

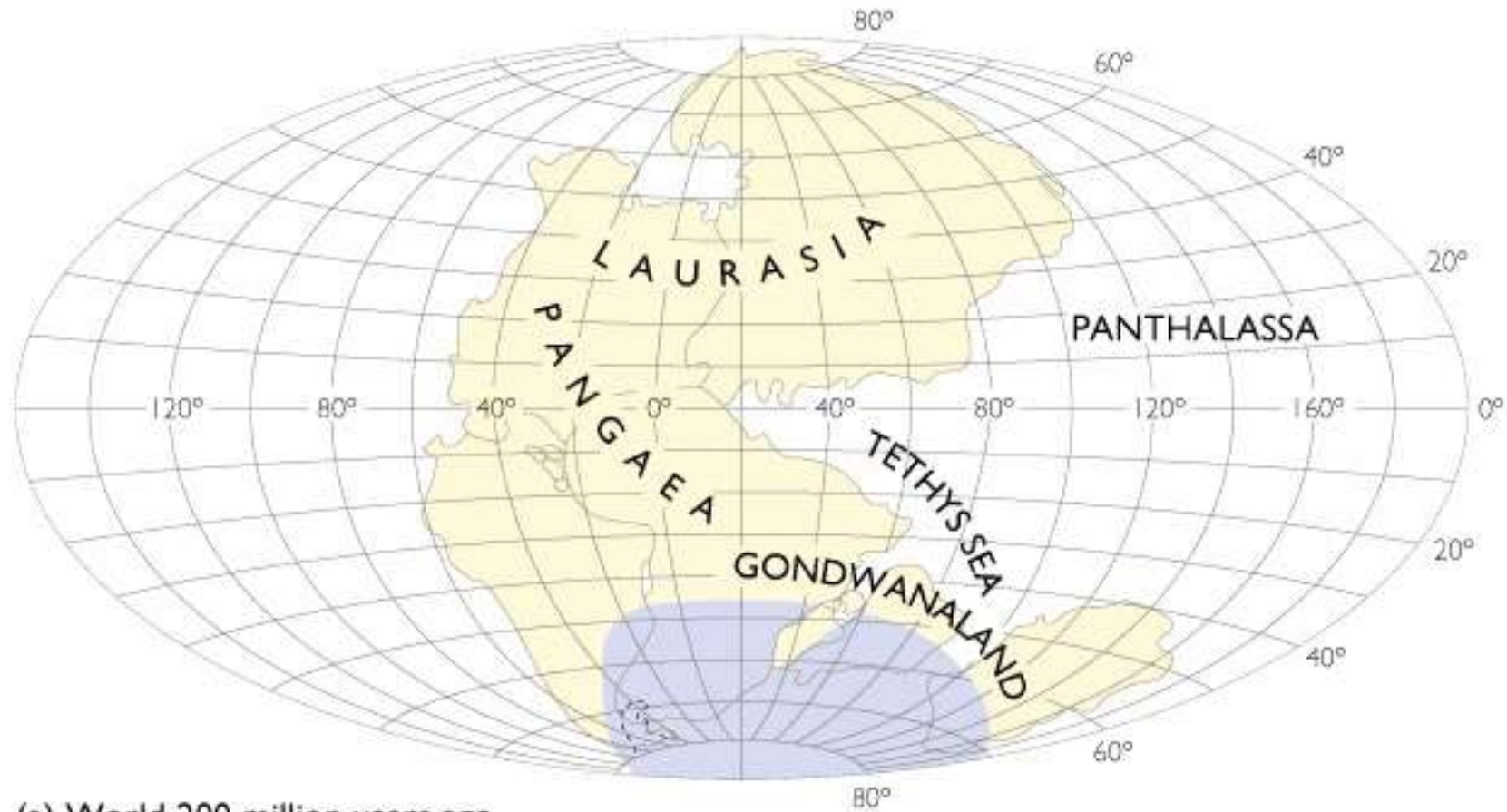


The South American Plate is moving northward away from a suspected impact crater in the South Atlantic





# Breakup of Pangaea



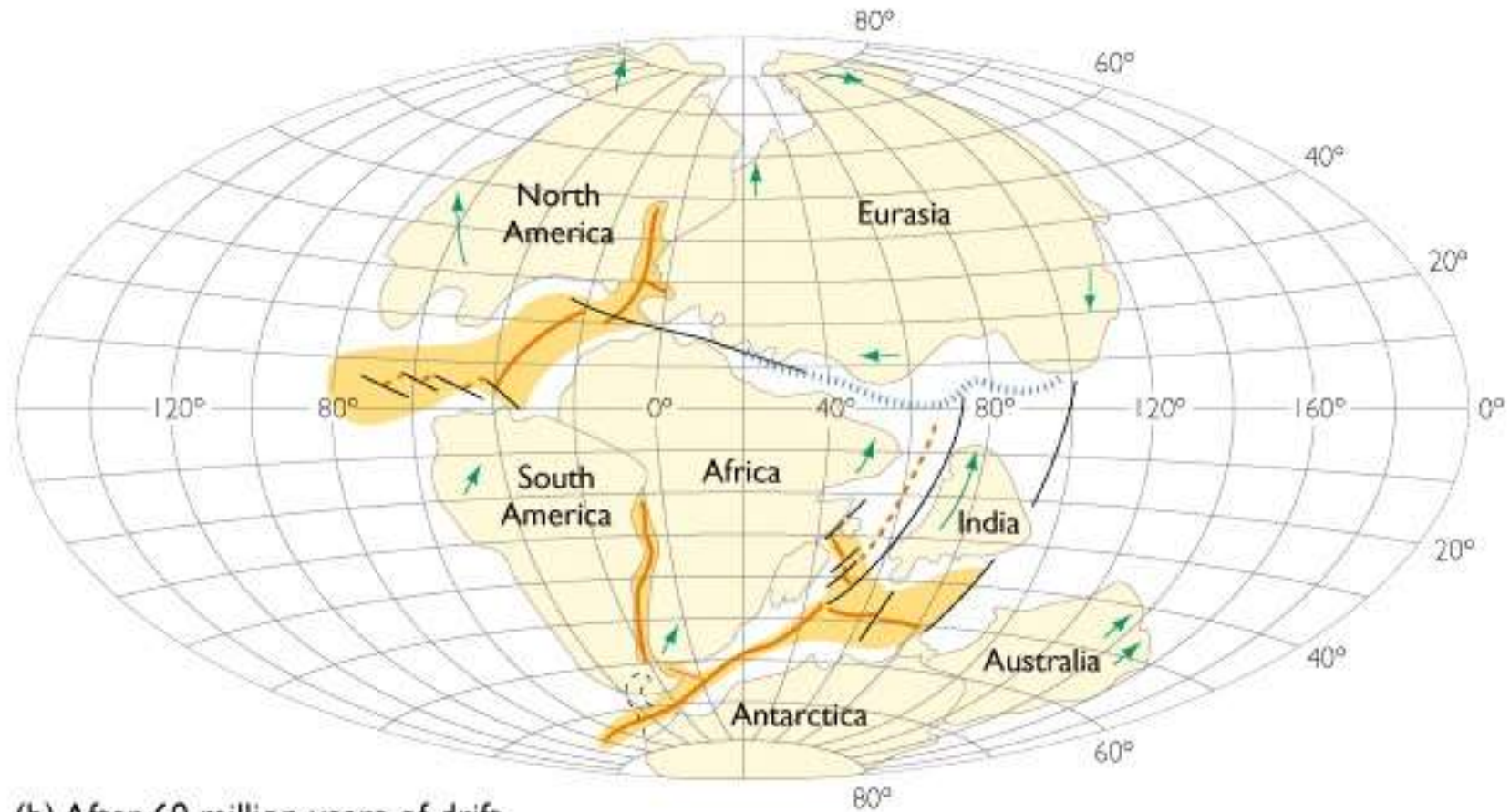
(a) World 200 million years ago

After Dietz & Holden, 1970

After Dietz & Holden, 1970

200 Ma

# Breakup of Pangaea



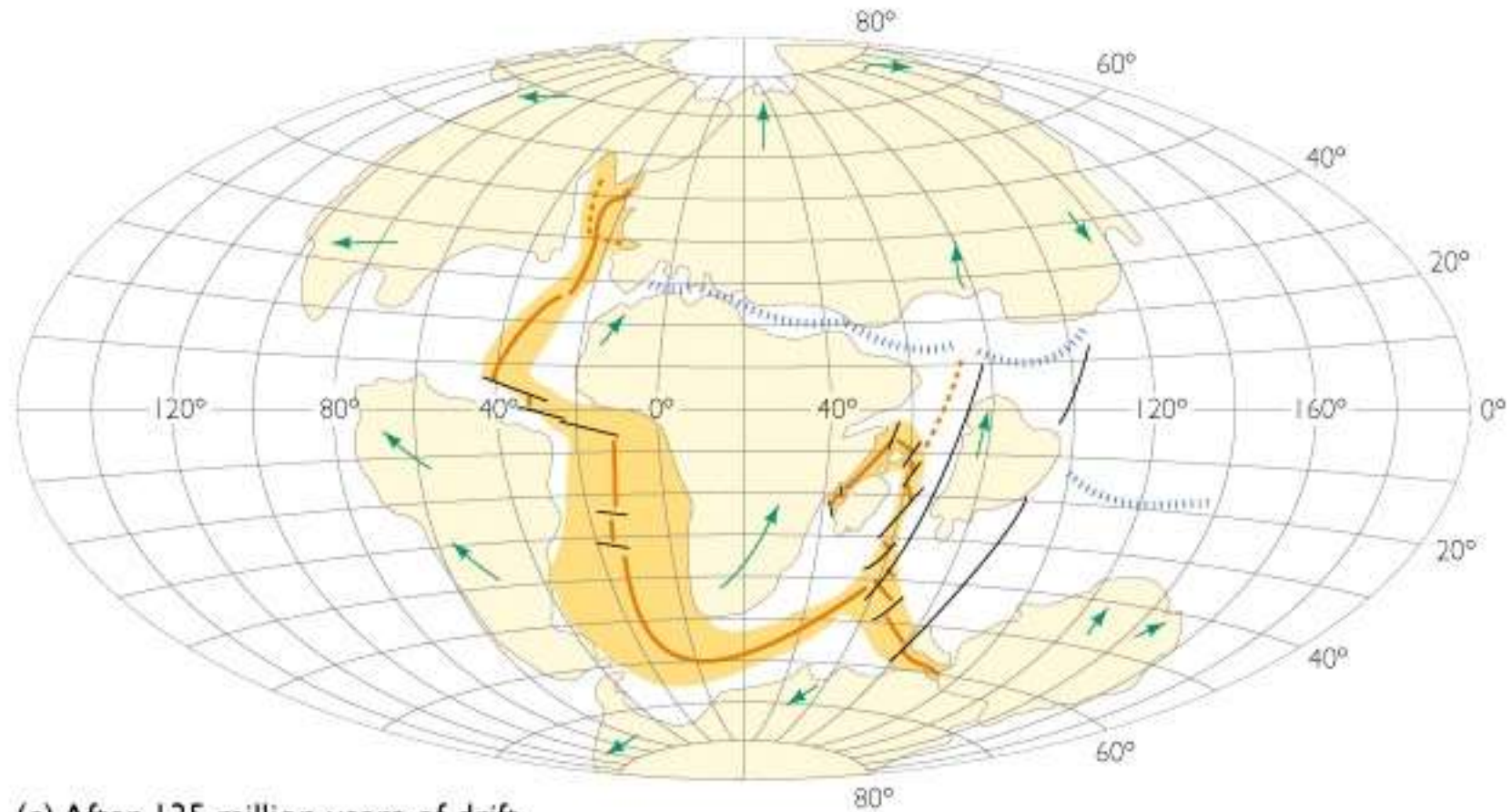
(b) After 60 million years of drift  
(140 million years ago)

After Dietz & Holden, 1970

After Dietz & Holden, 1970

140 Ma

# Breakup of Pangaea

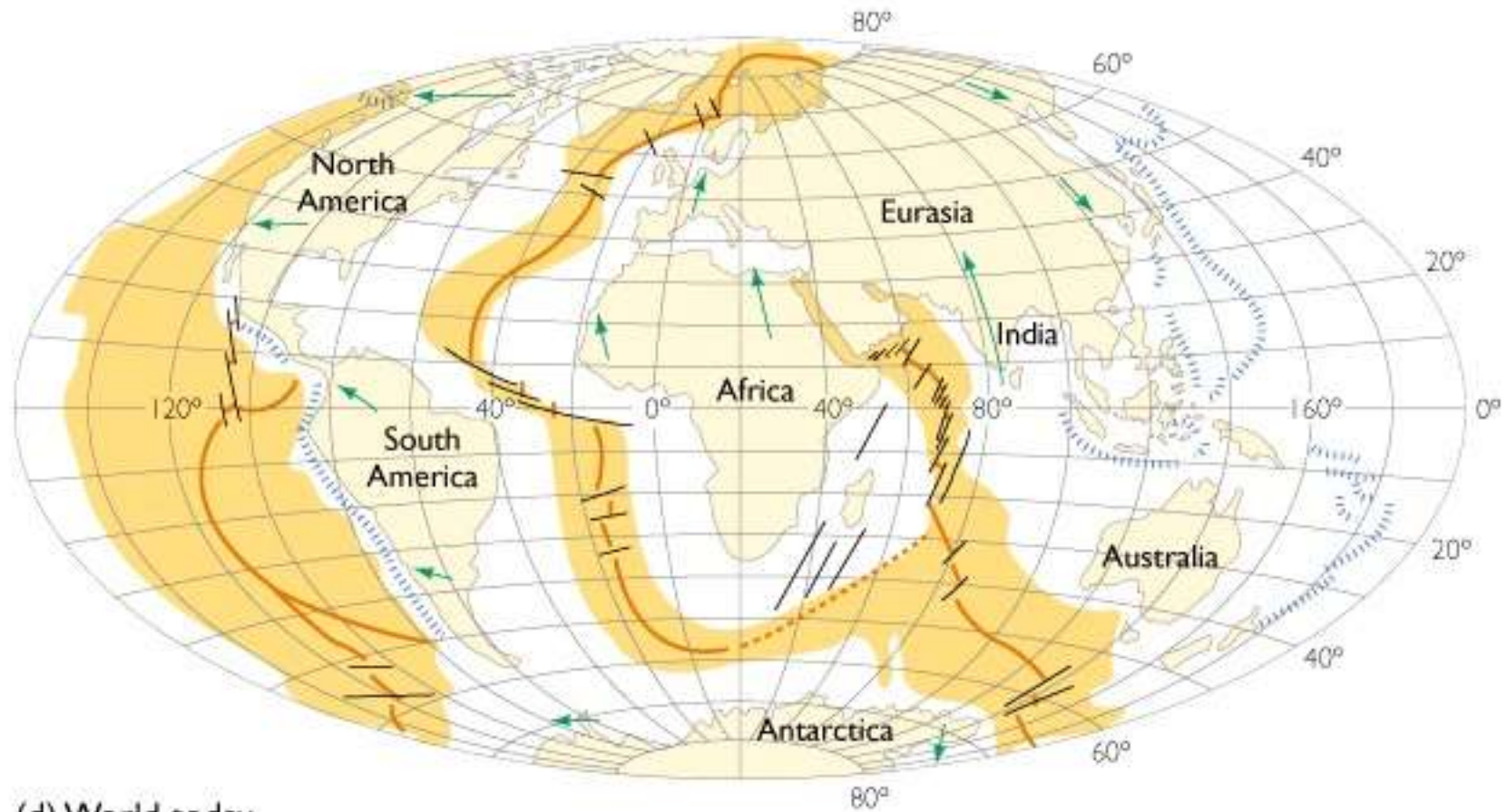


(c) After 135 million years of drift  
(65 million years ago)

After Dietz & Holden, 1970

65 Ma

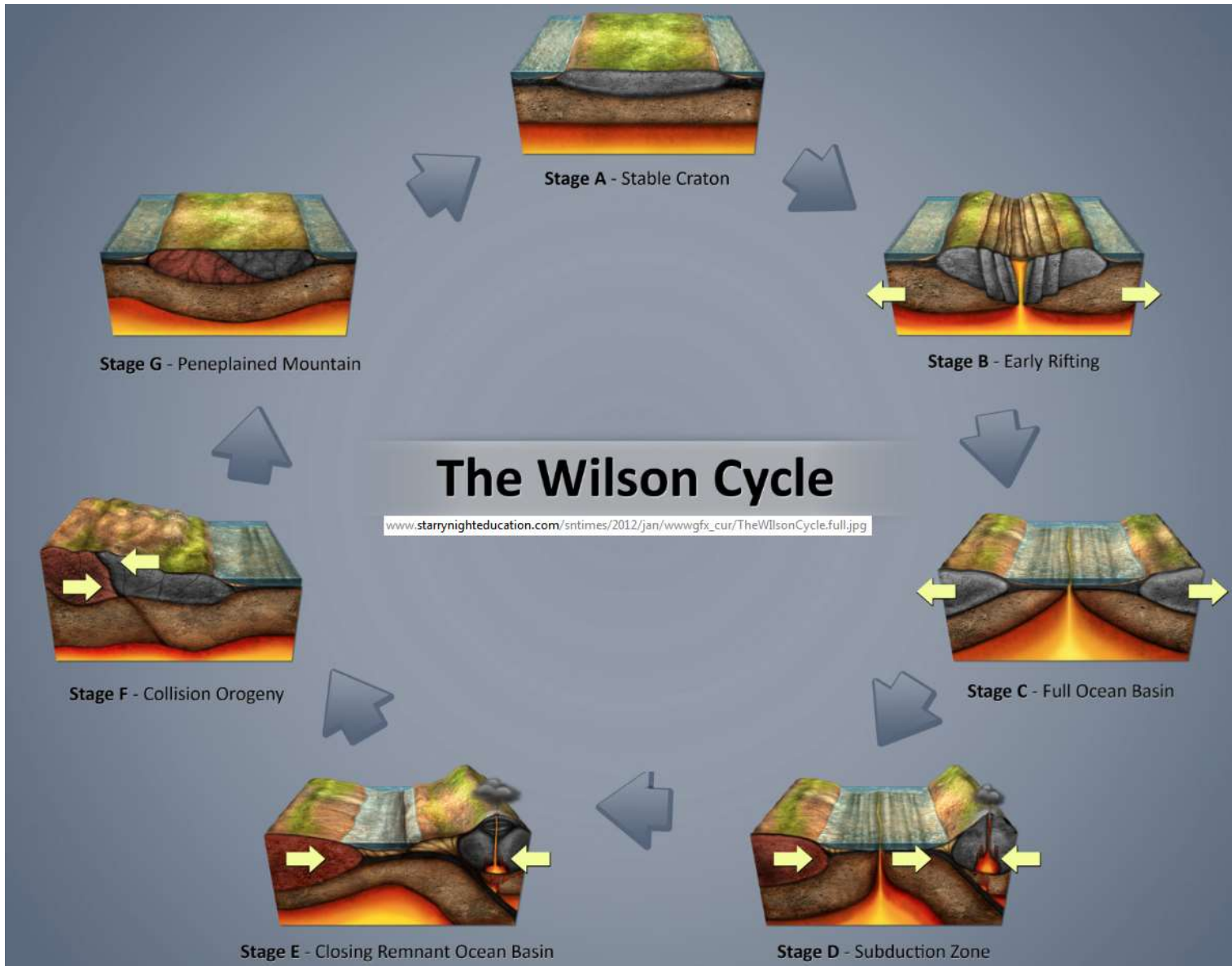
# Breakup of Pangaea



(d) World today

After Dietz & Holden, 1970

0 Ma



Geological data indicate that supercontinents like Pangea form, break up, and re-form in a cycle spanning about 500 million years.

The breakup forms rift valleys within the supercontinent that eventually becomes a long, linear sea as the crust is depressed below sea level.

As the width of the narrow sea continues to expand an open ocean develops and the continental masses drift apart.

For some reason, the continental masses aggregate anew, and the cycle begins again.